Do statistical models capture the dynamics of the magnetopause during sudden magnetospheric compressions?

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Key Points:

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15	• Measured magnetopause location is statistically closer to the Earth than Shue et
16	al. (1998) modelled for storm sudden commencements (SYM-H \geq 15 nT).
17	• When the magnetopause is compressed below 8 R_E , the average measured loca-
18	tion is > 1 R_E inside of the Shue et al. (1998) model location.
19	• Extreme magnetopause compressions rarely reach the outer radiation belt, there-
20	fore rapid outward radial transport is required to fully explain most shadowing
21	events.

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22 Abstract

Under periods of strong solar wind driving, the magnetopause can become compressed, 23 playing a significant role in draining electrons from the outer radiation belt. Also termed 24 'magnetopause shadowing', this loss process has traditionally been attributed to a com-25 bination of magnetospheric compression and outwards radial diffusion of electrons. How-26 ever, the drift paths of relativistic electrons and the location of the magnetopause are 27 usually calculated from statistical models and, as such, may not represent the time-varying 28 nature of this highly dynamic process. In this study, we construct a database $\sim 20,000$ 29 spacecraft crossings of the dayside magnetopause to quantify the accuracy of the com-30 monly used Shue et al. (1998) model. We find that, for the majority of events (74%), the 31 magnetopause model can be used to estimate magnetopause location to within $\pm 1 R_E$. 32 However, if the magnetopause is compressed below 8 R_E , the observed magnetopause 33 is greater than 1 R_E inside of the model location on average. The observed magnetopause 34 is also significantly displaced from the model location during storm sudden commence-35 ments, when measurements are on average 6% closer to the radiation belts, with a max-36 imum of 42 %. We find that the magnetopause is rarely close enough to the outer ra-37 diation belt to cause direct magnetopause shadowing, and hence rapid outward radial 38 transport of electrons is also required. We conclude that statistical magnetopause pa-39 rameterizations may not be appropriate during dynamic compressions. We suggest that 40 statistical models should be only be used during quiescent solar wind conditions, and sup-41 plemented by magnetopause observations wherever possible. 42

43 1 Introduction

Understanding the dynamics of the Van Allen radiation belts is a key challenge in 44 understanding the terrestrial space environment. The response of the radiation belts dur-45 ing geomagnetic storm-times is highly variable; storms may result in a net increase, a 46 net decrease, or indeed no net response at all (Reeves et al., 2003). At any given time, 47 a balance of acceleration and loss mechanisms could be contributing to the overall mor-48 phology of the radiation belts. It has been proposed that during storms there are repeat-49 able phases in radiation belt response; a net loss phase where electron losses dominate, 50 and a net acceleration phase where more electrons are accelerated than lost (Murphy et 51 al., 2018). Furthermore, the intensity of the net loss phase of a storm can control the level 52 of electron acceleration of electrons from the seed population in the following net-acceleration 53 phase (Bingham et al., 2018). Understanding the multiple sources of electron losses is 54 vital to understanding radiation belt dynamics during geomagnetic storms. 55

There are a variety of acceleration, transport and loss mechanisms that play a role 56 in shaping the radiation belt environment; from gyro-resonant interaction on kHz timescales 57 through to large-scale topological changes of the magnetosphere by solar wind-magnetosphere 58 interaction. Gyro-resonant wave-particle interactions between keV 'seed' electrons, in-59 jected into the inner magnetosphere during substorms, and Very Low Frequency (VLF) 60 whistler-mode waves, act to energize radiation belt electrons to MeV energies (Summers 61 et al., 1998; Horne & Thorne, 1998; Horne et al., 2005; Baker et al., 1998; Meredith et 62 al., 2002; Forsyth et al., 2016). Ultra Low Frequency (ULF) waves transport electrons 63 through radial diffusion (e.g. Fälthammar (1965); Jaynes et al. (2015)) and can play a 64 role in electron acceleration through drift-resonant wave-particle interactions (e.g., Elkington 65 et al. (1999); Mann et al. (2013)). Radial electron transport via ULF wave activity results in betatron acceleration (deceleration) of electrons as electrons are transported ra-67 dially inwards (outwards). Even without strong ULF wave activity, electrons may still 68 be adiabatically transported radially outwards if the ring current is enhanced, and drift-69 ing electrons will adiabatically decelerate (Dessler & Karplus, 1961; McIlwain, 1966)). 70

Whilst outwards transport and subsequent deceleration of electrons contribute to
 decrease in electron flux at a given energy, non-adiabatic effects account for significant

and irreversible loss of electrons from the radiation belts (Li et al., 1997; H.-J. Kim & 73 Chan, 1997). Loss mechanisms act to drain the radiation belts either into interplanetary 74 space or Earth's atmosphere. Again, gyro-resonant wave-particle interaction plays an es-75 sential role by depositing electrons into the atmospheric loss cone, through pitch-angle 76 scattering processes (Thorne & Kennel, 1971; Miyoshi et al., 2008; Gamble et al., 2008; 77 Ukhorskiv et al., 2010; Rodger et al., 2015). Localized, compressional ULF wave fields 78 may also play a role in precipitating relativistic electrons into the atmosphere (Rae et 79 al., 2018). Large scale topological changes to the geomagnetic field will also result in elec-80 tron loss if electron drift paths intersect the magnetopause (Li et al., 1997; Green et al., 81 2004; K. C. Kim et al., 2008; Saito et al., 2010). Such loss is not through precipitation 82 into the atmosphere, but rather loss from the magnetosphere itself, known as magnetopause 83 shadowing. The dayside magnetosphere may shrink due to compressions by enhanced 84 solar wind dynamic pressure and/or magnetopause erosion under southwards IMF (Gosling 85 et al., 1982; Sibeck et al., 1989; Dmitriev et al., 2014). Note that in this paper we use 86 the term compressed to synonymously refer to the inwards movement of the magnetopause 87 due to both pressure balance variations and magnetosphere erosion under southwards 88 IMF. We refer to two distinct types of magnetopause shadowing throughout this paper. 89 When the magnetopause is suddenly compressed within the outer radiation belt on time 90 scales similar to electron drift periods, such as during interplanetary shocks (Sibeck et 91 al., 1989), then electron drift paths directly intersect the magnetopause. We refer to this 92 as direct magnetopause shadowing. We distinguish 'indirect' magnetopause shadowing 93 as the combined process of outwards radial diffusion towards a compressed magnetopause. Hence, during indirect magnetopause shadowing the initial particle drift path does not 95 have to directly intersect the magnetopause boundary. Indirect magnetopause shadow-96 ing explains electron loss at comparatively low L shells where the magnetopause would 97 never directly impact (e.g. Brautigam and Albert (2000); Miyoshi et al. (2003); Y. Sh-98 prits et al. (2006); Loto'Aniu et al. (2010); Turner et al. (2012); Morley et al. (2010); Rodger 99 et al. (2019)) 100

The relative contributions of magnetopause shadowing and precipitation through-101 out a geomagnetic storm are not well understood. Previous work has shown that mag-102 netopause shadowing plays a clear role in electron flux drop out events (Y. Shprits et 103 al., 2006; Morley et al., 2010; Turner et al., 2012). Morley et al. (2010) studied 67 so-104 lar wind stream interface regions and showed electron flux decreased at L^* as low as 4 105 up to a day before the arrival of the stream interface at the bow shock. For these events, 106 the Shue et al. (1997) magnetopause model location reached a minimum of L = 8.5, which 107 is outside of where the losses were observed. Thus, Morley et al. (2010) attributed this 108 statistical loss to combined outward radial diffusion towards a compressed magnetopause. 109 Using the same event list, Hendry et al. (2012) analyzed precipitating electron flux mea-110 sured by the Polar Operational Environmental Satellites (POES). The authors observed 111 a large increase in precipitation following the arrival of the stream interface. During this 112 period of high electron precipitation, Morley et al. (2010) observed a net increase in elec-113 tron flux. Interestingly, Hendry et al. (2012) did not observe any increase in precipitat-114 ing electron flux during the electron flux drop out itself. It therefore appears that the 115 majority of losses prior to the stream interface arrival occur via magnetopause shadow-116 ing. 117

In order to understand the roles of direct or indirect shadowing on electron losses 118 observed by Morley et al. (2010), the position of the magnetopause and the last closed 119 drift shell (LCDS) needs to be known (Olifer et al., 2018). Both the magnetopause lo-120 cation and LCDS are calculated by models with a variety of underlying assumptions that 121 are likely violated during magnetopause compressions. For example, empirical magne-122 topause models (e.g. Shue et al. (1997, 1998)) assume the magnetopause is in an equi-123 librium position, and LCDS calculations assume that the magnetospheric field can be 124 accurately represented by global magnetic field models (e.g. Tsyganenko et al. (2003)). 125 Since we can measure the magnetopause location with relative accuracy as compared to 126

the LCDS, we choose to focus on how well a widely-used statistical magnetopause model 127 performs, with specific reference to dynamic times. We choose to analyze the Shue et 128 al. (1998) magnetopause model as it is widely used for radiation belt purposes (e.g. by 129 Morley et al. (2010); Loto'Aniu et al. (2010); Herrera et al. (2016); Olifer et al. (2018); 130 Turner et al. (2012); Murphy et al. (2015)). Previous statistical studies have shown the 131 Shue et al. (1998) model overestimated magnetopause location by $\sim 1 R_E$ at higher lat-132 itudes within the cusp region (Case & Wild, 2013). In this study we focus on the equa-133 torial subsolar point, where the LCDS is closest to the magnetopause. 134

135 In this study we construct a multi-spacecraft database of magnetopause crossings. We use this database to investigate the dynamics of the real magnetopause for events 136 which could lead to magnetopause shadowing and hence radiation belt loss events (Morley 137 et al., 2010). In order to do this, we first complete a statistical analysis of the measured 138 magnetopause location as compared to the Shue et al. (1998) model, identifying condi-139 tions under which the measured magnetopause location is significantly different to the 140 model, such as during interplanetary shocks and storm sudden commencements. We then 141 show how well our statistical results hold for a case study of the 2013 St. Patrick's day 142 storm, which is known to have a clear and well-studied radiation belt response (e.g. Albert 143 et al. (2018); Olifer et al. (2018); Ma et al. (2018)). Finally, we discuss whether a sta-144 tistical correction of the Shue et al. (1998) magnetopause model is useful in determin-145 ing the relative contributions of direct and indirect magnetopause shadowing during elec-146 tron dropout events. 147

¹⁴⁸ 2 The Shue et al. (1998) Magnetopause Model

Shue et al. (1997) carried out a best fit of a simple parabolic function to 553 mag-149 netopause crossings made by the ISEE 1 and 2, AMPTE/IRM and IMP 8 satellites. This 150 functional form depends only on the north-south component of the IMF and the solar 151 wind dynamic pressure, D_p , which determine the subsolar standoff distance and tail flar-152 ing angle of the parabola. The measurements of the magnetopause used to fit the model 153 were taken during solar wind conditions in the range 0.5 nPa $< D_p < 8.5$ nPa and IMF 154 $-18 \text{ nT} < B_z < 15 \text{ nT}$. The authors discuss that the fitted model does not give real-155 istic values of tail flaring angle for IMF B_z and D_p outside of these ranges. Shue et al. 156 (1998) refitted the functional form of the Shue et al. (1997) model to include the non-157 linear dependence of dynamic pressure, D_p on tail flaring angle, and also the impact of 158 IMF B_z on subsolar standoff distance. The revised Shue et al. (1998) model gives a much 159 improved representation of the magnetopause during values of D_p and B_z in their range 160 of fitting data. As the Shue et al. (1998) model is easily implemented, it is extensively 161 used to estimate magnetopause standoff distance in radiation belt physics. For brevity, 162 we henceforth refer to this model as the 'Sh98' model. The Sh98 model has frequently 163 been applied to understanding electron flux dropout events, where magnetopause shad-164 owing contributes to global radiation belt electron loss (Morley et al., 2010; Loto'Aniu 165 et al., 2010; Herrera et al., 2016; Olifer et al., 2018). 166

It must be noted that the Sh98 model assumes a rigid parabolic magnetopause that 167 is in equilibrium with cylindrical symmetry around the aberrated Sun - Earth line. This 168 implies that the magnetopause responds instantaneously and globally to any changes in 169 upstream solar wind conditions. In reality, the magnetopause is much more dynamic. For 170 example, surface waves are driven at the magnetopause which oscillates about its equi-171 librium (Plaschke, Glassmeier, Sibeck, et al., 2009). Cahill and Winckler (1992) also ob-172 served large solar wind compressions which break equilibrium and drive magnetopause 173 oscillation. In addition, the magnetopause does not have cylindrical symmetry: Case and 174 Wild (2013) completed a statistical comparison of the Sh98 model to a database of high 175 latitude Cluster magnetopause crossings, demonstrating that Sh98 model tended to over-176 estimate the standoff distance by $\sim 1 R_E$ near the cusps. 177

Shue et al. (1998) discuss the uncertainty arising from magnetopause motion. The 178 authors calculate uncertainty as a function of IMF B_z , D_p , and solar-zenith angle. The 179 authors concluded that any deviation from the modelled average position due to, for ex-180 ample, magnetopause oscillations, are represented by the known Sh98 model uncertain-181 ties. Using the method described in Shue et al. (1998), Figure 1 presents how the Sh98 182 uncertainties vary with solar-zenith angle and IMF B_z orientation, given (a) moderate 183 (IMF $|B_z|$ and D_p of 4 nT and 2 nPa respectively), and (b) strong (IMF $|B_z|$ and D_p 184 of 15 nT and 8 nPa respectively) solar wind driving. Figure 1 (a) shows that uncertainty 185 increases from $\sim 0.15 \text{ R}_E$ to 0.3 R_E for both northward and southward IMF $|B_z|$ and 186 is \sim 0.025 R_E higher for southward IMF as compared to northward. The same trends 187 of increasing uncertainty with solar-zenith angle is true for higher solar wind driving (Fig-188 ure 1 (b)), but there is a larger difference between southward and northward orientated 189 IMF, and under these conditions southward IMF now has a lower uncertainty than un-190 der northward IMF conditions. Comparing the uncertainties for southward IMF across 191 moderate and higher solar wind driving (across Figures 1 (a) and (b)), it is interesting 192 to note that the uncertainty is lower for higher solar wind driving across all solar-zenith 193 angles. In contrast, for northward IMF, the uncertainties are increased. In this study, 194 we take the maximum uncertainty in the modelled subsolar magnetopause standoff dis-195 tance to be $\sim 0.2 \text{ R}_E$, and $\lesssim 0.4 \text{ R}_E$ across the modelled dayside magnetopause (solar-196 zenith angles less than or equal to 90° , which effectively corresponds to the entirety of 197 the dayside magnetosphere). 198

¹⁹⁹ **3** Dataset and Methodology

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3.1 Building a database of magnetopause observations

In order to compare the Sh98 model with measurements of the real magnetopause, 201 we compile a database of spacecraft crossings of this boundary. When a spacecraft crosses 202 the magnetopause, the measured magnetic field will transition between a strong, steady 203 and northwards orientated field within the magnetosphere, and a rapidly varying mag-204 netic field characteristic of the magnetosheath, that may be orientated in any direction. 205 Plasma density transitions from low values in the outer magnetosphere, to higher den-206 sities in the magnetosheath where the shocked solar wind piles up and stagnates (Crooker 207 & Siscoe, 1975). 208

To conduct our analysis, we have created a new database of magnetopause crossings which is further supplemented by databases from three previous studies, which are described in Table 1. These databases used automated or semi-automated classification methods. For full details of the automated algorithms and data sets we refer the reader to Plaschke, Glassmeier, Sibeck, et al. (2009); Case and Wild (2013); Raymer (2018).

Satellite	# Crossings	Timespan	Authors
Geotail Mukai et al. (1994); Kokubun et al. (1994)	8,548	1996 - 2015	Raymer (2018)
THEMIS Auster et al. (n.d.); McFadden et al. (2008)	6,697	2007	Plaschke, Glassmeier, Sibeck, et al. (2009)
Cluster Balogh et al. (2001)	2,688	2002 - 2010	Case and Wild (2013)

Table 1. Details of three existing databases of spacecraft magnetopause crossings.



Figure 1. Uncertainty of the Shue et al. (1998) magnetopause model as a function of solarzenith angle for southward and northward orientated IMF magnitude (a) $|B_z| = 4$ nT and $D_p = 2$ nPa; (b) $|B_z| = 15$ nT and $D_p = 8$ nPa. These calculations make use of uncertainty calculation described in the original Shue et al. (1998) model paper.

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To take full advantage of the THEMIS dataset since its launch in 2007, we developed a semi-automated method to classify an additional 12,621 THEMIS magnetopause crossings from 2007 to 2016. Data from THEMIS probes A, D and E were used from 2007 - 2016 and THEMIS B and C from 2007 - 2010, after which these spacecraft were moved to lunar orbit (Angelopoulos, 2010). Only spacecraft crossings of the dayside magnetopause $(X_{GSE} > 0R_E)$ were classified as this is the region electron drift paths are most likely to intersect the magnetopause.

Our approach classifies magnetopause crossings from the THEMIS Flux gate magnetometer (FGM) (Auster et al., n.d.) and Electrostatic Analyzer (ESA) (McFadden et al., 2008) instruments. We first created an algorithm which used a set of criteria to classify possible magnetopause crossing candidates, these crossing candidates were then manually verified on a daily basis. Data where missclassifications are clearly within the magnetosheath or magnetosphere were discarded, although a small number of missclassifications may still exist due to human error.

To create the crossing criteria we manually classified 18 magnetopause crossings by the THEMIS E probe between 17:00 and 23:00 UT on the 16th June 2007. The crossing criteria were then empirically determined by optimizing the number of these crosspirically determined as follows: When THEMIS crossing from the magnetosphere to the magnetosheath:

- ²³⁵ 1. The change in the B_z component of the magnetic field, in GSM coordinates, must ²³⁶ be less than -0.6 nT s⁻¹, and the change in ion density must be greater than 0.08 ²³⁷ cm⁻³ s⁻¹;
- 238 2. Within the magnetosphere, the average B_z component of the magnetic field must 239 be greater than 5 nT and the average ion density must be less than 7 cm⁻³ for 240 a 48 s interval;
- 3. The first two crossing criteria must be met within a 60 s interval.

If THEMIS is crossing from the magnetosheath to the magnetosphere, we reverse 242 the first criteria. To prevent spurious measurements from high frequency noise when cal-243 culating the first criteria, we down-sampled measurements of the B_z component of FGM 244 measurements from a 3 s resolution to 24 s and ESA measurements of ion density ware 245 reduced from 3 s to 36 s resolution. Once these crossings were visually verified, the database 246 contained 34,428 confirmed magnetopause crossings. We have removed multiple cross-247 ings of the magnetopause that occurred within 10 minutes, retaining only the innermost 248 crossing for each probe. The innermost crossing was used so that our database is com-249 parable to the Sh98 model, which used only the innermost crossing in a series of cross-250 ings to fit the model. Removing multiple crossings reduced the database to 12,621 cross-251 ings. 252

The Plaschke, Glassmeier, Sibeck, et al. (2009) magnetopause database also con-253 tains a large number of multiple magnetopause crossings due to the nature of their study 254 of magnetopause oscillations. Multiple crossings within 10 minutes are also removed from 255 this database, retaining only the innermost crossing for each probe. Finally we cross ref-256 erenced the Plaschke, Glassmeier, Sibeck, et al. (2009) database with our THEMIS database 257 to ensure THEMIS crossings are not double counted. As before, the innermost crossing 258 of the magnetopause from either database within a 10 minute interval was retained. This 259 reduces the Plaschke, Glassmeier, Sibeck, et al. (2009) database to 1,910 crossings and 260 the database we classified for this study is reduced to 11,821 crossings. 261

This renders a final database of 24,967 THEMIS, Cluster and Geotail magnetopause 262 crossings spanning almost two solar cycles from 1996 - 2016. Figure 2 shows the spatial 263 distribution of magnetopause crossings over all solar-zenith angles for $2 \times 2 R_E$ bins. Fig-264 ure 2 shows the number of crossings on the dayside magnetopause, with the maximum 265 number of crossings in any bin is 1,892 crossings between 8 to 10 $R_E X_{GSM}$ and 0 to 266 $-2 R_E Z_{GSM}$ (panel (c)). The lowest number of magnetopause crossings occur on the 267 magnetopause tail ($X_{GSM} < 0$ R_E in Figure 2 (a and c)) where many spatial bins only 268 contain a single crossing. The coverage of the down-tail magnetopause is significantly 269 less than the dayside since these crossings are taken only from the Geotail database. Note, 270 in the following analysis, we take all magnetopause measurements from the dayside mag-271 netopause only (from 06-18 MLT) since our main focus is to investigate the role of mag-272 netopause shadowing on the radiation belts. This reduces our database to a total of 19,973 273 measurements of the dayside magnetopause, which we use to perform our statistical anal-274 ysis for the remainder of this study. 275



Figure 2. 2D histogram of magnetopause crossings in the (a) X_{GSM} - Y_{GSM} plane; (b) Y_{GSM} - Z_{GSM} plane; (c) X_{GSM} - Z_{GSM} plane.

We use solar wind data provided by the NASA/Goddard Space Flight Centers OMNI dataset through Coordinated Data Analysis Web (CDAWeb; https://omniweb.gsfc.nasa.gov/) that comprises solar wind measurements from the ACE, Wind, IMP 8 and Geotail missions. The solar wind data is propagated to the bow shock nose and has a temporal resolution of 5 minutes. It is expected that propagation time from the bow shock to the magnetopause is similar to this 5 minute resolution (Villante et al., 2004).

We also use the Symmetric Horizontal (SYM-H) index at a 5 minute resolution, as the de-facto high-resolution version of the Dst index (Wanliss & Showalter, 2006). The SYM-H index is calculated in a similar manner to Dst by ground based, mid-latitude magnetometer stations. This data is also provided in the OMNI dataset.

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3.2 Comparing magnetopause observations to a modelled location

²⁸⁷ We define ΔR as the radial distance between the measured location of a spacecraft ²⁸⁸ magnetopause crossing, R_{SC} , and the distance to the abberated Sh98 model magnetopause, ²⁸⁹ R_{Mod} , for the same solar-zenith angle of the spacecraft, such that $\Delta R = R_{Mod} - R_{SC}$. ²⁹⁰ If $\Delta R > 0$ then the model overestimates the magnetopause location, i.e., the Sh98 mag-²⁹¹ netopause is located at a larger radial distance than the measured magnetopause. Con-²⁹² versely, if $\Delta R < 0$ then the Sh98 model underestimates the magnetopause location, i.e., the Sh98 model is closer to the Earth than the measurement. Finally, if $\Delta R = 0$ to within an uncertainty of ± 0.4 R_E, then we conclude that the model and the measurement agree.

It is also important to estimate the position of the subsolar magnetopause where an electron drift path is more likely to intersect the magnetopause. By assuming that the functional shape of the Sh98 magnetopause is correct (i.e., that the shape and flaring angle, α , is correct) then we can project spacecraft measurements from any dayside magnetopause crossing to the abberated subsolar point, $R_{0_{SC}}$, by rearranging the Sh98 functional form (Plaschke, Glassmeier, Sibeck, et al., 2009; Plaschke, Glassmeier, Auster, et al., 2009);

$$R_{0_{SC}} = R_{SC} \left(\frac{2}{1 + \cos\theta}\right)^{-\alpha} \tag{1}$$

where θ is the solar-zenith angle of the spacecraft crossing position, calculated by taking the inverse cosine of the dot product between the aberrated Sun-Earth line and the position vector of the spacecraft in GSE coordinates. We then define the difference between the modelled subsolar standoff distance and the measured equivalent subsolar standoff distance as $\Delta R_0 = R_{0_{Mod}} - R_{0_{SC}}$, where $R_{0_{Mod}}$ is the modelled subsolar standoff distance.

Finally, we also define the percentage change in distance to be $\Delta R/R_{SC}$ to normalize for times where there is a compressed or expanded magnetopause, and in order to compare crossings across all dayside solar-zenith angles to each other.

312 4 Results

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4.1 Statistical evaluation of magnetopause location

The distributions of ΔR and ΔR_0 are shown in Figure 3 (a) and (b), respectively. 314 Figure 3 (a) shows ΔR to not be normally distributed as the mean and median values 315 are not equal; the mean $\Delta R = 0.13 \text{ R}_E$ and the median $\Delta R = 0.05 \text{ R}_E$. This asymmetric 316 try implies that there are a higher number of instances where the measured magnetopause 317 is closer to Earth than the modelled distance. Furthermore, 74 % of measurements lie 318 within one standard deviation of the mean, which in this case is $0.97 R_E$. The upper and 319 lower quartiles of ΔR are - 0.43 R_E and 0.64 R_E respectively. The difference between 320 the median and the mean is less than the Sh98 model uncertainty of $\pm 0.4 R_E$, but there 321 is a large spread in ΔR , with only 40 % of measurements being within $\leq 0.4 R_E$. 322

Figure 3 (b) shows the ΔR_0 is also not a normal distribution as the mean and median values are not equal; with a mean $\Delta R_0 = 0.09 \text{ R}_E$ and median $\Delta R_0 = 0.05 \text{ R}_E$. Furthermore, 70 % of measurements occur within a standard deviation of the mean, where $\sigma = \pm 0.84 \text{ R}_E$. The upper and lower quartiles of ΔR_0 are = - 0.40 R_E to 0.56 R_E respectively. The difference between the median and the mean is less than the Sh98 model uncertainty of $\pm 0.2 \text{ R}_E$ 1, but there is a large spread in ΔR_0 , with only 24 % of measurements being within $\leq 0.2 \text{ R}_E$.

To see how ΔR varies for different measured standoff distances, in Figure 4 we in-330 vestigate the median magnetopause distance calculated by the Sh98 model, R_{Mod} , as a 331 function of experimentally observed magnetopause distance, R_{SC} . We note this Figure 332 describes spacecraft crossings at all measured solar-zenith angles, R_{SC} , rather than equiv-333 alent subsolar standoff, R_0 . The shaded area shown in the figure indicates the inter-quartile 334 range of R_{SC} measurements. Within the shaded region, it can be seen the distribution 335 is closest to the line of unity, so median $R_{Mod} \simeq R_{SC}$ indicating that the Sh98 model 336 is accurately calculating magnetopause standoff distance at locations between 10.6 and 337 12 R_E . However, there is clearly a different gradient than unity. A multiple linear re-338



Figure 3. (a) The distribution of ΔR for measurements of the dayside magnetopause. (b) The distribution of ΔR_0 for measurements of the dayside magnetopause. The solid blue line shows the median value for each panel and the dotted blue lines show the inter-quartile range.

gression to the distribution of median R_{Mod} is given by the purple line in Figure 4. We find that the experimentally measured magnetopause distance as a function of median modelled magnetopause distance is best described by $R_{SC} = \frac{R_{Mod} - 3.68}{0.68}$.

In order to assess whether different solar wind conditions are influencing these un-342 der and overestimations of magnetopause location by the Sh98 model, we examine dis-343 tributions of $\Delta R/R_{SC}$ for varying solar wind dynamic pressure and north-south IMF. 344 These figures are included in supplementary material 1 and 2 respectively. Whilst there 345 was a weak relationship between $\Delta R/R_{SC}$ and D_p , there was no evidence that strong 346 dynamic pressures $(D_p > 4 \text{ nPa})$ are associated with large positive $\Delta R/R_{SC}$. Similarly, 347 $\Delta R/R_{SC}$ showed a tendency to increase when IMF B_z magnitude increased, but this 348 was not true across all B_z magnitudes. 349

We further examine the distribution of $\Delta R/R_{SC}$ for varying geomagnetic condi-350 tions. Figure 5 (a) shows a 2D histogram of $\Delta R/R_{SC}$ as a function of the SYM-H in-351 dex. We column normalize the distributions, since there are many more measurements 352 during geomagnetically quiet times (-50 nT \lesssim SYM-H \lesssim 15 nT) than for the rest of the 353 distribution. Figure 5 (a) demonstrates that the relationship between $\Delta R/R_{SC}$ and the 354 geomagnetic conditions, as defined by SYM-H index, varies depending on phase of the 355 geomagnetic storm. For quiet times (SYM-H between -50 and 15 nT), the maximum oc-356 currence probabilities are peaked and centred on zero. However, for geomagnetic storm-357 time conditions (SYM-H \leq -50 nT), the median offset between measurement and model 358 varies greatly between - 5 % to 15 % for decreasing SYM-H. Moreover, for positive SYM-359 H, $\Delta R/R_{SC}$ has a near-constant positive offset that increases with increasingly positive 360 SYM-H. This positive offset indicates that the magnetopause is closer to Earth than the 361 model prediction. We note that large positive and sudden increases in SYM-H typically 362 correspond to the storm sudden commencement phase (SSC) of a geomagnetic storm. 363 During the SSC, SYM-H index can increase by 10s of nT on minute timescales (Dessler et al., 1960) in response to the arrival of an interplanetary shock front rapidly compress-365 ing the dayside magnetosphere. Figure 5 (a) would therefore suggest that the magne-366 topause is closer to Earth by up to 15 % during SSC. 367

To further demonstrate the observed distribution of offsets in Figure 5 (a) between measurements and the Sh98 model observed, we examine the distribution of $\Delta R/R_{SC}$



Figure 4. Purple diamonds show the median standoff distance calculated by the Shue et al. (1998) model, R_{Mod} , corresponding to spacecraft magnetopause crossing measured at a given standoff distance, R_{SC} . The error bars show the propagated error of the Shue et al. (1998) model (see Section 2). The blue line gives where $\overline{R}_{Mod} = R_{SC}$. The shaded area indicates the interquartile range (10.6 to 12.0 R_E) of observed magnetopause distance, R_{SC} .

for geomagnetically quiet times (SYM-H between -50 and 15 nT), during the main phase 370 of geomagnetic storm (SYM-H \leq -50 nT) and for storm sudden commencement (SYM-371 $H \ge 15$ nT), shown in Figure 5 (b). There are 19,140 measurements of $\Delta R/R_{SC}$ for SYM-372 H between -50 and 15 nT. This distribution is peaked at $\Delta R/R_{SC} = 0$ % with upper 373 and lower quartiles of - 4 % and 6 % respectively. Thus, during relatively quiet times 374 the observed location of the magnetopause is as located inside the Sh98 model location 375 as often as it is located outside the model location. In contrast, when SYM-H \leq -50 nT 376 or SYM-H \geq 15 nT, the peak of the distribution is positive; 2 % and 4 % respectively, 377 with upper and lower quartiles of -3 % and 9 % for SYM-H \leq -50 nT and 0 % and 10 378 % for SYM-H \geq 15 nT. There are 601 magnetopause measurements during SYM-H \leq 379 -50 nT and 137 measurements for SYM-H \geq 15 nT. We use the MannWhitney U test 380 (Nachar, 2008) to confirm that the SSC and main storm phase distributions are statis-381 tically different as compared to the quiet time distribution, to a 95 % confidence level. 382 As such, during storm times (SYMH \leq -50 nT) it is more likely that the magnetopause 383 will be inside of the model location. During periods when SYMH ≥ 15 nT, which typ-384 ically correspond to SSCs, the magnetopause location is almost exclusively inside of the 385 model location. Thus, the magnetopause is statistically closer to the Earth than the Sh98 386 model during both the main phase of a geomagnetic storm and during storm sudden com-387 mencement. 388

To test the more extreme deviations from the Sh98 model, we perform a superposed epoch analysis (SEA) of solar wind drivers during strongly positive SYM-H conditions.



Figure 5. (a) Column normalized distribution of percentage change in magnetopause standoff distance $(\Delta R/R_{SC})$ as a function of SYM-H. Column medians are indicated by black crosses. (b) Probability distributions of $\Delta R/R_{SC}$ separated by geomagnetic conditions; quiet times are shown by the grey histogram (-50 nT < SYM-H < 15 nT), storm sudden commencement phase is shown by the blue histogram (SYM-H \geq 15 nT), main storm phase is shown by the purple histogram (SYM-H \leq -50 nT).

We select events for this analysis where there is a peak in SYM-H which exceeds 15 nT, 301 where there is a spacecraft measurement of the magnetopause within a day of the peak 392 SYM-H. Epoch time zero, t_0 , is chosen as the peak value of SYM-H. We then perform 393 the superposed epoch analysis for ± 1 day of t_0 . Figure 6 shows the results of this SEA. 394 In total there were 392 individual events used in the analysis, and 3,629 spacecraft cross-395 ings of the magnetopause across all of the epochs used. Figure 6 (f) shows median $\Delta R/R_{SC}$ 396 at a 2 hour resolution for the superposed epochs, whereas Figures 6 (a - e) have a 5 minute 397 resolution. The 2 hour resolution of Figure 6 (f) was chosen such that the variability of 398 $\Delta R/RSC$ through the epoch analysis is clear, whilst maximizing the number of crossings used to calculate each median value through the epoch analysis. 400

Figure 6 shows strong evidence of solar wind discontinuities at t_0 characteristic of forwards interplanetary shocks; a sudden increase in temperature and an increase in magnetic field strength following t_0 (Figure 6 (a) and (b)), and a sharp peak in D_p at t_0 (Figure 6 (c)). It is well understood that fast forwards interplanetary shocks play a large role in the storm sudden commencement phase due to enhancement of magnetopause currents (e.g. Taylor (1969)). In particular empirical relationships have been derived between SSC amplitude and the change in the square root of D_p at the shock/discontinuity (Russell et al., 1992).

In response to the sudden dynamic pressure increase, the Sh98 model demonstrates 409 a compression of the median subsolar magnetopause from 10.7 R_E to 8.7 R_E (Figure 6 410 (e)). We observe that the SYM-H index shows a tendency to become negative follow-411 ing t_0 in Figure 6 (d). Further investigation showed that 33 % of the epochs contained 412 moderate to intense geomagnetic storms with minimum SYM-H \leq -50 nT. A further 30 413 % of epochs contained a minimum of SYM-H between -30 and -50 nT, indicating weak 414 geomagnetic storms (Loewe & Prölss, 1997). This supports our suggestion that a peak 415 in SYM-H \geq 15 nT indicates a storm sudden commencement phase of shock-driven ge-416 omagnetic storms. 417



Figure 6. A superposed epoch analysis of (a) interplanetary magnetic field strength B (blue) and B_z (purple); (b) solar wind temperature, T; (c) solar wind dynamic pressure D_P ; (d) SYM-H index; (e) subsolar standoff distance of the Shue et al. (1998) magnetopause model, R_{0Mod} ; (f) Percentage difference in radial distance between measured location of the magnetopause and the Shue et al. (1998) magnetopause model, $\Delta R/R_{SC}$. Epoch time zero is defined as the time that SYM-H reaches a peak ≥ 15 nT. The purple lines show median values and the inter-quartile range is denoted by the shaded regions. The vertical dotted line shows t_0 .

The median percentage difference between the spacecraft measurements of the mag-418 netopause and the Sh98 model, $\Delta R/R_{SC}$, is noted to be relatively small and slowly vary-419 ing between - 2 % and 2 % until 4 hours (0.2 days) before $t_0 \Delta R/R_{SC}$ rapidly increased 420 to 6% (Figure 6 (f)). At the same time, the upper quartile of SYM-H exceeds 0 nT (Fig-421 ure 6 (d). Following this rapid increase, $\Delta R/R_{SC}$ reached a maximum of 6 % at t_0 . Me-422 dian values of $\Delta R/R_{SC}$ remain high until 2 hours (0.1 days) after t_0 and, as shown in 423 6 (f), the entire inter-quartile range is greater than 0%, which means that in the ma-424 jority of cases the Sh98 model is overestimating magnetopause distance. At times greater 425 than 2 hours after t_0 , median $\Delta R/R_{SC}$ decreases but remains positive, fluctuating be-426 tween 0 % and 3 %, though the inter-quartile range is notably larger than times preced-427 ing $t_0 + 2$ hours. 428

It is important to comment that in Figure 5 (b) the median $\Delta R/R_{SC}$ was calcu-429 lated as 4 % when we used a threshold of SYM-H \geq 15 nT to define magnetopause mea-430 surements taken during a SSC. Whereas in the SEA presented in Figure 6 (f), median 431 $\Delta R/R_{SC} = 6$ % at t_0 , which is defined as the time SYM-H peaks at a value greater than 432 15 nT. This difference is because the SEA of $\Delta R/R_{SC}$ has a resolution of 2 hours: Mag-433 netopause measurements which occur within an hour of the SYM-H peak ≥ 15 nT are 434 included in the median calculation, though SYM-H may be less than 15 nT at the time 435 of the crossing. 436

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4.2 Case study: 2013 St Patrick's Day Storm

Finally, we inspect a case study of a large geomagnetic storm associated with a sig-438 nificant radiation belt response. The chosen event is the 2013 St. Patricks day storm, 439 which has both a magnetopause crossing when SYM-H > 15 nT and is a large geomag-440 netic storm where magnetopause shadowing should have played an important role in ra-441 diation belt electron losses. The 2013 St. Patricks day storm has been studied extensively 442 by the Geospace Environment Modeling (GEM) program radiation belt modelling fo-443 cus group Quantitative Assessment of Radiation Belt Modeling (QARBM) as a radia-444 tion belt modeling event to quantitatively assess (e.g. Albert et al. (2018); Olifer et al. 445 (2018); Ma et al. (2018)). This event is a CME-driven geomagnetic storm in which a rapid 446 flux dropout of the outer radiation belt precedes strong enhancements in electron flux 447 during the geomagnetic storm (Olifer et al., 2018). 448

The 2013 St. Patrick's day storm has 93 individual magnetopause crossings observed by the THEMIS probes in our database between 14-20 March, all of which we have visually verified. From Figure 7 there are three separate solar wind pressure enhancements that lead to magnetopause variations on the 14, 15 and 17 March 2013, indicated by vertical dashed lines. We discuss each of these pressure enhancements in turn.

At around 13:00 UT on the 14 March 2013, there was a small increase in D_p of up to 5 nPa (Figure 7 (a)), which had a small corresponding SYM-H increase to 16 nT but no obvious radiation belt response (Figure 7 (b), (d)). The model Sh98 standoff location was compressed to 8.8 R_E. The equivalent standoff distance of magnetopause crossings during this compression, $R_{0_{SC}}$, agree remarkably well with the Sh98 location (Figure 7 (c)).

Around 06:00 UT on the 15 March 2013, there was a second comparatively small 460 increase in D_n of up to 6 nPa (Figure 7 (a)). In this case there was a clear and rapid in-461 crease in SYM-H by 20 nT, to a peak of 25 nT. There was a reduction in the $\sim 1 \text{ MeV}$ 462 electron fluxes by a factor of ~ 10 at the outer boundary of the radiation belt, for L > 463 $6 R_E$ (Figure 7 (d)), which persisted until the major geomagnetic disturbance which be-464 gan on the 17 March 2013. For the 15 March pressure pulse, the Sh98 model standoff 465 location was compressed to 8.3 R_E . The observed magnetopause crossings continued to 466 agree these model values until the magnetopause is maximally compressed at 07:00 UT. Immediately following this maximum compression on the 15th there are two magnetopause 468



Figure 7.

March 2013 St. Patrick's day storm case study from 14-20 March 2013; (a) solar wind dynamic pressure, D_p ; (b) SYM-H index, with the blue and purple horizontal lines denoting SYM-H = -15 and -50 nT, respectively; (c) the subsolar standoff distance of the Sh98 model in blue, black crosses denoting magnetopause crossing distance at any solar zenith angle, and purple crosses denoting equivalent standoff distance of those magnetopause crossings, $R_{0_{SC}}$; (d) ~ 1 MeV electron flux from the Van Allen Probes A and B MagEIS instruments to illustrate radiation belt activity. Vertical dashed lines indicate pressure enhancements. Blue and purple shaded areas denote the SSC and main phase of the geomagnetic storm respectively. A zoomed-in plot of panels (a) - (d) during the geomagnetic storm on 17 March is also shown.

⁴⁶⁹ crossings where the observed magnetopause is $0.9 R_E$ and $1.2 R_E$ (10 % and 15 %) closer ⁴⁷⁰ to the radiation belts than the Sh98 model distance.

On 17 March 2013, the CME arrival was accompanied by a sudden increase in D_p 471 from 1 nPa to 15 nPa. The SYM-H index responded accordingly, with a sharp increase 472 to 31 nT, before the main phase of the storm where SYM-H decreased down to -131 nT. 473 During the main phase of the storm, the ~ 1 MeV electron fluxes decreased by around 474 2 orders of magnitude, a reduction that persisted for 7 hours. Enhancements resulting 475 from the storm dominated over the losses on the 18 March 2013; on this day the $\sim 1 \text{ MeV}$ 476 electron fluxes increased by 3 orders of magnitude and the radial peak in flux moved to 477 lower L of ~ 4 R_E. The model Sh98 standoff location (R_{Mod}) was compressed in response 478 to the pressure enhancement, and was as close to the Earth as $6.1 R_E$ during the main 479 phase of the storm. At 08:45 UT, where the Sh98 model output was at it's minimum stand-480 off distance, the subsolar projection of an observed THEMIS E magnetopause crossing 481 was 5.7 R_E, which is 0.4 R_E (7 %) closer to Earth than the Sh98 model calculation of 482 $6.1 R_E$. During the storm sudden commencement, there was one crossing of the mag-483 netopause made by THEMIS D at 06:48 UT, with an equivalent subsolar standoff dis-10/ tance of 6.4 R_E. At this time, the Sh98 model was calculated as 7.3 R_E, a difference of 485 $0.9 R_E$ (or 14 %) closer to Earth than the model calculation. Equivalent subsolar stand-486 off measurements during the main phase of the storm were perhaps even more variable, 487 ranging between 5.7 and 10 R_E , indicating that the Sh98 model does not reflect the true 488 magnetopause location during this highly disturbed time. Taking the model uncertainty 489 as $\sim 0.2 \text{ R}_E$ at the subsolar point (following the calculations presented in Figure 1), only 490 15~% of measurements on the 17 March 2013 were within this error. The Sh98 model 491 underestimated standoff distance by > 0.2 R_E for 40 % of measurements, and overes-492 timated standoff distance by > 0.2 R_E for 45 % of measurements. 493

⁴⁹⁴ 5 Discussion and Conclusions

The ability to accurately calculate the magnetopause standoff distance is integral to the process of modelling and prediction of trapped electron fluxes in the outer radiation belt. An accurate magnetopause location is central to accurately determining whether radiation belt losses will occur via direct magnetopause shadowing, indirect magnetopause shadowing, or not at all.

Olifer et al. (2018) studied a series of geomagentic storms, where a model magne-500 topause and last closed drift shell (LCDS) could be determined. These authors also used 501 the Shue et al. (1998) magnetopause model and concluded that there was a strong cor-502 respondence between the variation in the LCDS and measured electron fluxes during these 503 case studies. Olifer et al. (2018) concluded that their results implied that indirect mag-504 netopause shadowing, i.e. outward radial transport combined with enhanced ULF wave 505 radial diffusion, played a key role in relativistic losses during rapid flux dropout events. 506 Albert et al. (2018) investigated the behaviours of different LCDS models, finding that 507 models of the LCDS differ distinctly in L* depending on the assumptions used, the dif-508 ferent magnetic field model inputs, and calculation procedures. Interestingly, Olifer et 509 al. (2018) calculated that the Sh98 magnetopause model was, at times, Earthwards of 510 the LCDS prior to storm sudden commencement (SSC) (see lower panel of Figure 3, Olifer 511 et al. (2018)). That the LCDS can lie outside of the Sh98 model location exemplifies that 512 LCDS models should be used with caution. Matsumura et al. (2011) used an empirical 513 outer boundary of the radiation belt as a proxy for the last closed drift shell. These au-514 thors found that this empirical boundary is well correlated with the magnetopause stand-515 off distance as calculated by Shue et al. (1997) during loss events, when the outer bound-516 ary of the radiation belt moved Earthwards in conjunction with a compression of the Shue 517 et al. (1997) model. Given that we are unable to measure the LCDS, but we can mea-518 sure the magnetopause location, we have tested the validity of the most common mag-519 netopause model used for radiation belt physics. 520

In this study, we constructed an empirical database of $\sim 20,000$ spacecraft crossings of the dayside magnetopause. We compared the locations of each crossing with the

predicted Sh98 model location given the prevailing solar wind conditions, provided by 523 the OMNI database. The radial difference between the measured and predicted magne-524 topause location was distributed about zero, with upper and lower quartiles of \sim - 0.5 525 R_E and 0.6 R_E respectively, for all dayside locations (Figure 3 (a)) and when mapped 526 to the subsolar point (Figure 3 (b)). However, the distributions were slightly skewed to-527 wards positive values for both ΔR and ΔR_0 as the means of both distributions were ~ 528 0.1 R_E with a standard deviation of ~ 1 R_E . This means that the Sh98 model accurately 529 represented the magnetopause location to within $\sim 1 R_E$, on average. Figure 7 corrob-530 orates this finding, as the Sh98 model and the measured magnetopause are in agreement 531 during the 14 - 15 March 2013 time period. It must be noted that the calculation of R_{0sc} 532 and ΔR_0 assumes that the shape of the Sh98 model, specifically the level of tail flaring, 533 α , is correct. If a spacecraft crossing is at a large solar-zenith angle (i.e. not near the mag-534 netopause nose) this method of mapping to the subsolar point may introduce error in 535 $R_{0_{SC}}$ or ΔR_0 calculations if α is inaccurate. 536

Further, whilst we found that the predicted Sh98 model magnetopause location was 537 accurate to within $\sim 1 R_E$ of the observed magnetopause locations between 10.5 R_E and 538 12 R_E , the uncertainty increased for more extreme cases, i.e. when the measured mag-539 netopause location was outside of this range (Figure 4). On average, the Sh98 model un-540 derestimated standoff distance for crossings measured at distances > 12 R_E , and over-541 estimated standoff distance for crossings measured at distances $< 10.6 R_E$. We applied 542 a multiple linear regression to the observed and average modelled values and found that 543 across all prevailing conditions between 1996 - 2016, the relations can be described by a linear function $R_{SC} = \frac{R_{Mod} - 3.68}{0.68}$. This fit of the model to our crossing database may 544 545 suggest that a simple correction made to the Sh98 location would better reflect the av-546 erage measured location. However, we emphasize that the linear regression shown in Fig-547 ure 4 should not be used to correct the Sh98 model on an event by event basis without 548 careful consideration. This is particularly important for values of R_{Mod} smaller than those 549 used in the linear regression $(R_{Mod} < 7.4 \text{ R}_E)$, where the prediction of R_{SC} for the lin-550 ear regression becomes unrealistically small. For example, for a modelled prediction of 551 $6.6 R_E$, the linear regression would imply that the magnetopause position would be 3.9 552 \mathbf{R}_E . 553

Discrepancies between measurements and the model for large observed magnetopause 554 distances $(R_{SC} > 12.0 \text{ R}_E)$ could be due, in part, to inaccuracies in the paraboloid Sh98 555 model shape, i.e., the magnetopause is closer than the model near the nose, and further 556 away near the flanks, which would be suggestive of a more flared magnetotail. Further 557 inaccuracies in the paraboloid Sh98 model shape may arise from the no-axisymmetric 558 shape of the magnetopause, i.e. dawn-dusk asymmetries (Haaland et al., 2017) and in-559 dentations due to the magnetospheric cusp regions (Case & Wild, 2013). We also con-560 sidered whether the difference between measurements and the model for small observed 561 magnetopause distances ($R_{SC} < 10.6 \text{ R}_E$) could be due to the Sh98 inaccurately rep-562 resenting the influence of dynamic pressure or IMF on the magnetopause location (Sup-563 plementary Information 1 and 2 respectively). Whilst dynamic pressure and IMF do not 564 appear to be responsible for systematic discrepancies between measured magnetopause 565 location and the Sh98 model, we would recommend that the Sh98 model should only be 566 used in the range of 0.5 nPa $< D_p < 8$ nPa and -15 nT $< B_z < 10$ nT. This is based 567 on the distribution of median $\Delta R/R_{SC}$ measurements in Supplementary Figures S1 and 568 S2 respectively, and the range of dynamic pressures and IMF magnitudes for which Shue 569 et al. (1998) had magnetopause measurements to fit the Sh98 model (Section 2). Finally, 570 we note that the observed discrepancy between model and measurements may be, in part. 571 due to rapid solar wind fluctuations. Processes such as solar wind fluctuations would mean 572 that the magnetopause location is not in equilibrium, as assumed by the Sh98 model. 573 In this study we have shown that during dynamic times such as interplanetary shocks, 574 an average location will not reflect the true magnetopause location. Hence any empir-575 ical relationship should therefore be used with extreme caution. 576

We have show that the distance between the measured magnetopause and the mod-577 elled location varies for different geomagnetic conditions (Figure 5). We highlight that, 578 for increasingly positive SYM-H, the magnetopause location is increasingly overestimated 579 by the Sh98 model. This overestimate may be up to a maximum median of 13~% between 580 40 nT \leq SYM-H \leq 60 nT, and maximum single event value of 42 % at a SYM-H of 18 581 nT (Figure 5 (a)). We identify these periods of positive SYM-H as the storm sudden com-582 mencement (SSC) phase (Figure 6). Hence, for increasingly large SSCs, the magnetopause 583 location can be expected to be significantly closer to the Earth than previously thought. 584 Figure 6 shows a superposed epoch analysis (SEA) of solar wind drivers during strongly 585 positive SYM-H conditions associated with SSCs. We find that the driver of strong pos-586 itive increases of magnetopause compressions show characteristics of fast forward shocks. 587 The strong positive increases in SYM-H were found to be associated with magnetospheric 588 compressions (Figure 6 (e)). At the maximum SYM-H, the magnetosphere was maxi-589 mally compressed and observations of the magnetopause were overestimated by 6 % on 590 average by the Sh98 magnetopause model (Figure 6 (f)). 591

Solar wind pressure pulses and fast forward shocks have been known to have an 592 associated radiation belt response (e.g. Sibeck et al. (1989); Hietala et al. (2014); Kilpua 593 et al. (2019)), which is usually attributed to shock driven ULF waves which radially dif-594 fuse electrons towards the magnetopause (e.g. Claudepierre et al. (2010)). In particu-595 lar, relativistic electron flux in the outer radiation belt has been observed to drop out 596 in response to a stream interface of high speed solar wind streams; Morley et al. (2010) 597 showed results of a SEA where electron flux drops out at L^* as low as 4.5 in response 598 to high speed solar wind stream interface regions. The authors observed that the mag-599 netopause standoff distance becomes compressed to 8.5 R_E and concluded that electron 600 losses occurred by more indirect magnetopause shadowing i.e., magnetopause compres-601 sion and rapid outward radial transport. The results we have presented in Figure 6 for 602 fast forward shocks, such as high speed solar wind stream interfaces, would suggest that 603 it is highly likely that the magnetopause is compressed significantly closer to the outer 604 radiation belt than the Sh98 model calculates. In Figure 8 we investigate using a cor-605 rection to the Sh98 magnetopause model for the 67 stream interface events identified by 606 Morley et al. (2010). For each individual epoch, we identify the maximum value of SYM-607 H. Then, for ± 12 hours (0.5 days) from this peak in SYM-H, we increase or decrease 608 the Sh98 standoff distance by a factor that is time-dependent according to the $\Delta R/R_{SC}$ 609 results shown in Figure 6 (f), e.g. for t_0+5 hours of a SYM-H peak, R_{Mod} is decreased 610 by 5 %. Figure 8 (a) shows a SEA of SYM-H during the SI events and Figure 8 (b) shows 611 a SEA of the Sh98 subsolar magnetopause standoff (pink-purple colours) and corrected 612 magnetopause standoff distances are shown by blue colours. 613

Figure 8 shows a SEA of R_{Mod} (pink-purple colours) and R_{Cor} is shown by blue 614 colours. In addition to the inter-quartile range of the Sh98 modelled magnetopause po-615 sition during the SIs, which reached a minimum of 8 R_E , the full range of values was as 616 low as 6 R_E . Given that the Sh98 model standoff distances had values within geosta-617 tionary orbit (6.6 R_E) we find that, at least in some circumstances, that direct magne-618 topause shadowing may occur following a number of these SIs. Moreover, when we ap-619 ply a correction to the modelled standoff distance, we find that the estimated median 620 magnetopause location is compressed to 8.2 $\mathrm{R}_{E},$ with a lower quartile value of 7.6 $\mathrm{R}_{E},$ 621 and the minimum magnetopause compression during all the epochs was 5.9 R_E . If a me-622 dian magnetopause location is used then direct magnetopause shadowing would not be 623 predicted, regardless of whether a correction is applied to the Sh98 model or not. How-624 ever, direct magnetopause shadowing may still occur during more extreme conditions, 625 particularly during the SSC period. Figure 8 illustrated that, under more extreme or vari-626 able conditions, this standoff distance can be significantly closer to the Earth than the 627 Sh98 model. Given that our maximum difference between measurement and model is 42 628 % closer to the Earth during positive SYM-H, this would lead to the magnetopause be-629 ing well inside geostationary orbit and as close as 5 R_E . We suggest that this may hap-630

Morley_cor_new.png

Figure 8. Superposed epoch analysis of 67 high speed solar wind stream interface events identified by Morley et al. (2010). (a) The dark purple line shows median SYM-H index and the light purple shaded region shows the inter-quartile range; (b) The dark purple line shows the median Shue et al. (1998) subsolar standoff distance of the magnetopause, R_{Mod} , the light purple shaded area shows the inter-quartile range and the light pink line shows the minimum standoff distance of R_{Mod} at a given epoch time. The dark blue line shows the median corrected magnetopause standoff distance, the shaded blue area shows the interquartile range, and the light blue line shows the minimum standoff distance of R_{Cor} . The correction factor is based on variations in $\Delta R/R_{SC}$ associated with a peak in SYM-H index (Figure 6 (f))

pen during dynamic periods such as SSC but that, during the main and recovery phase
 of storms, it is more likely that outwards radial diffusion must still be invoked to explain
 electron losses i.e., via indirect magnetopause shadowing.

In order to investigate the time-dependent accuracy of model magnetopause motion during a radiation belt dropout event, we studied the 2013 St Patrick's day storm. We found that, during more quiescent times before the geomagnetic storm, the observed and model magnetopause locations are very similar between the 14-16 March 2013. However, at the end of the 16 March 2013 and 17 March 2013 the Sh98 model magnetopause standoff distance was rarely accurate compared to observed magnetopause crossings. 85 % of observed magnetopause standoff distances were either greater than (40 %) or less

than (45 %) the Sh98 magnetopause standoff by distance greater than the model uncer-641 tainty. From our measurements of the magnetopause we calculated the equivalent day-642 side magnetopause to reach a minimum of 5.7 R_E , 0.4 R_E closer to the outer radiation 643 belt than the Sh98 calculation. In panel (d) of Figure 7, we observed that this compres-644 sion of the magnetopause will have been capable of causing direct shadowing of the outer 645 radiation belt. Indirect shadowing will also have played a role in this dropout event as 646 ULF wave power was high during this period (Ma et al., 2018), transporting electrons 647 at lower L shells towards the compressed magnetopause. Betatron deceleration of elec-648 trons as they are transported radially outwards could further contribute towards the ap-649 parent decrease in electron flux of a given energy channel. The combined result was that 650 the entire outer radiation belt decreased in flux by 2 orders of magnitude. 651

We now discuss several additional aspects concerning how our analysis might be 652 affected by both small-scale transitory structures in the magnetopause and by large-scale 653 motion of the magnetopause. Firstly, models such as the Sh98 model aim to character-654 ize the global shape and location of the magnetopause, but in reality the magnetopause 655 contains smaller scale structures. For example Kelvin-Helmholtz waves occur at the mag-656 netopause flanks due to an instability created by a velocity shear at the magnetopause 657 boundary layer (e.g. Pu and Kivelson (1983); Hasegawa et al. (2004)). Hot flow anoma-658 lies in the solar wind are known to decrease pressure in regions of the magnetosheath 659 for short periods of time (~ 7 mins) allowing the magnetopause to bulge outwards by 660 up to 5 R_E near the hot flow anomaly core (Sibeck et al., 1999; Jacobsen et al., 2009; 661 Archer et al., 2014). Conversely, fast magnetosheath jets can produce local magnetopause 662 indentations of up to $\sim 1-2 R_E$ depth if a jet penetrates to the magnetopause (Shue et 663 al., 2009; Amata et al., 2011; Hietala et al., 2014; Plaschke et al., 2016). Surface waves on the magnetopause have also been observed as a result of impinging magnetosheath 665 jets (Plaschke, Glassmeier, Sibeck, et al., 2009; Amata et al., 2011). If a magnetopause 666 crossing takes place in a location where the magnetopause is locally perturbed, then the 667 crossing may not represent the global magnetopause location, if such a thing exists. Not 668 only do these structures add uncertainty to the estimation of magnetopause location, they 669 potentially have effects on the dynamics of magnetospheric plasma. Both Kelvin-Helmholtz 670 waves and magnetospheric jets are known drivers of ULF waves (Southwood, 1974; Chen 671 & Hasegawa, 1974; Hughes, 1994; Claudepierre et al., 2008; Archer et al., 2013), which 672 act to diffuse magnetospheric plasma. Earthwards perturbations of the magnetopause 673 due to a fast magnetosheath jet near the subsolar point may intersect radiation belt elec-674 tron drift paths. What is more, local magnetopause compressions due to fast magnetosheath jets only occur for tens of seconds up to 3 minutes (Archer et al. 2012). If a magnetosheath 676 jet is sustained for minute time scales near the subsolar magnetopause, it could certainly 677 contribute towards a substantial loss of the ultra-relativistic electron population, which 678 have drift periods of ~ 5 minutes. However, electron losses in the outer radiation belt 679 have not yet been observed directly in connection with magnetosheath jets (Plashke et 680 al., 2018). We expect more global changes in magnetopause location to largely govern 681 total radiation belt dropout events as most of the relativistic electron population have 682 drift orbits longer than the time scale of a magnetosheath jet. 683

Secondly, in our analysis we use only the innermost of a sequence of magnetopause 684 crossings to represent the position of the magnetopause at that time. Measured mag-685 netopause crossings will primarily be due to the magnetopause passing over a quasi-stationary 686 spacecraft, and hence the minimum magnetopause location will lie somewhere inside the 687 spacecraft location. In part, this is addressed by the Shue et al. (1998) model, whereby 688 the innermost magnetopause crossing was taken to be the minimum standoff distance 689 in their model. However, during a large compression by an interplanetary shock, or lo-690 cal compression due to a fast magnetosheath jet, this would not reflect the minimum mag-691 netopause location. Moreover, any interplanetary shock that leads to an SSC will set the 692 magnetopause in motion until it reaches an equilibrium position, and so an average mag-693 netopause correction is not necessarily representative of specific event behaviour (Freeman 694

et al., 1995). Freeman et al. (1995) studied magnetopause motion during time varying 695 solar wind conditions, such as those studied in this paper. The authors found that, to 696 a first order approximation, the magnetopause behaves like a 2D elastic membrane and 697 exhibits oscillation of a damped harmonic oscillator in response to changes in solar wind 698 dynamic pressure. In their idealized system, the magnetopause oscillation is highly damped 699 with a natural eigenperiod of ~ 7 minutes. Hence, it is certainly possible that electrons 700 with drift periods of ~ 5 minutes could intersect the oscillating magnetopause location 701 when the magnetopause is undergoing this damped harmonic motion before settling to 702 a more equilibrium position. This would involve the total loss of ultrarelativistic elec-703 trons but only a small disturbance to the medium energy radiation belt electron pop-704 ulation - much like the reports of ultrarelativistic electron losses currently attributed to 705 EMIC wave-driven precipitation (e.g., Y. Y. Shprits et al. (2017); Aseev et al. (2017)). 706

Ideally, continuous observations of the magnetopause location would elucidate the 707 time-dependent response of the magnetopause to variable solar wind driving and geo-708 magnetic storms. These observations could be conducted by the Solar Wind - Magne-709 tosphere - Ionosphere Link Explorer, or 'SMILE', a small class science mission which is 710 under development between the European Space Agency and Chinese Academy of Sci-711 ences (Raab et al., 2016). This novel experiment will use observations of soft X-ray emis-712 sions from charge exchange interactions in the Earth's magnetosheath, from which a three-713 dimensional magnetopause location can be inferred. The SMILE mission provides a unique 714 opportunity to investigate the role of the global magnetopause on radiation belt dynam-715 ics. 716

5.1 Summary:

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- During periods of slowly varying solar wind conditions, and quiescent geomagnetic activity, we have found that the Sh98 magnetopause model is a good estimate of magnetopause location within $\pm 1 \text{ R}_E$.
- We highlight that the time-dependent response of the magnetopause to fast changes is solar wind conditions (e.g. interplanetary shocks) cannot be captured by a statistical magnetopause model such as the Shue et al. (1998) model. During such times, other parameterizations of the magnetopause location should be considered, supplemented by measurements of the magnetopause wherever possible.
- The time-dependent nature of the magnetopause must be taken into account for
 any realistic description of radiation belt electron losses through the magnetopause.
 In particular, we show that a new parameterization may be critical when quantifying electron flux dropouts in the radiation belts, particularly at very high energies.

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All data is publicly available via http://cdaweb.gsfc.nasa.gov.

The new THEMIS magnetopause database classified for this study is provided in the supplementary information file 'THEMIScrossings.txt'. THIS DATA WILL BE UP-LOADED TO A REPOSITORY BEFORE PUBLISHING; the supplementary file provided is only intended for reviewer reference. The repository to archive this data will be chosen from the AGU recommended website https://repositoryfinder.datacite.org/. F.A.S. was supported by a Science and Technology Funding Council (STFC) studentship. I.J.R. is supported in part by STFC grants ST/N000722/1 and ST/S000240/1,
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