- 1 Long-Lasting Geomagnetically Induced Currents and Harmonic Distortion
- <sup>2</sup> observed in New Zealand during the 07-08 September 2017 Disturbed Period
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- 16 Main point # 1: Analysis of a transformer in New Zealand, shows a sequence of large
- 17 Geomagnetically Induced Currents (GIC) associated with a storm period.
- 18 Main point # 2: Unique combination of measurements show primarily even harmonics
- 19 generated by transformer saturation when GIC >15 A.
- 20 Main point # 3: During study period only longer-lasting GIC generated observable
- 21 harmonics, but limited GIC impact from impulsive solar wind shock events.
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Abstract. Several periods of Geomagnetically Induced Currents (GIC) were detected in the 24 Halfway Bush substation in Dunedin, South Island, New Zealand, as a result of intense 25 geomagnetic storm activity during 06 to 09 September 2017. Unprecedented data coverage 26 from a unique combination of instrumentation is analyzed, i.e., measurements of GIC on the 27 single phase bank transformer T4 located within the substation, nearby magnetic field 28 perturbation measurements, very low frequency (VLF) wideband measurements detecting 29 the presence of power system harmonics, and high-voltage harmonic distortion 30 measurements. Two solar wind shocks occurred within 25 hours, generating four distinct 31 periods of GIC. Two of the GIC events were associated with the arrival of the shocks 32 themselves. These generated large but short-lived GIC effects that resulted in no observable 33 harmonic generation. Nearby and more distant magnetometers showed good agreement in 34 measuring these global-scale magnetic field perturbations. However, two subsequent 35 longer-lasting GIC periods, up to 30 minutes in duration, generated harmonics detected by 36 the VLF receiver systems, when GIC levels continuously exceeded 15 A in T4. Nearby and 37 more distant magnetometers showed differences in their measurements of the magnetic field 38 perturbations at these times, suggesting the influence of small-scale ionospheric current 39 structures close to Dunedin. VLF receiver systems picked up harmonics from the substation, 40 up to the 30th harmonic, consistent with observed high-voltage increases in even harmonic 41 distortion, along with small decreases in odd harmonic distortion. 42

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#### 44 **1. Introduction**

Rapid fluctuations of the Earth's magnetic field can lead to Geomagnetically Induced 45 Currents (GIC) flowing in high-voltage transformers and power systems [Pirjola & Boteler, 46 (2017)]. Fluctuating ionospheric current systems induce quasi-direct current (DC) in the 47 surface of the Earth, which can enter into man-made transmissions lines [Molinski, 2002]. 48 The effects of GIC on long transmission lines, such as the early telegraph systems, have been 49 noted since the 1840's [Boteler et al., 1998]. Typically the largest GIC levels occur for the 50 fastest rate of change of the geomagnetic field, with the most extreme changes being seen in 51 the region of the auroral electrojet [Birkeland, 1908; Cummings & Dessler, 1967]. However 52 the latitude of the electrojet varies depending on the level of geomagnetic activity [see 53 Oughton et al., 2017 for a comprehensive discussion], with high geomagnetic storming 54 displacing the electrojet equatorwards from geomagnetic high latitudes to mid-latitudes 55 [Thomson et al., 2011], which would include the southern regions of New Zealand's South 56 Island relevant to this study. 57

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During a GIC event unidirectional DC flux adds to the alternating current flux in 59 transformer cores for alternate half cycles potentially leading to peak magnetic flux levels 60 that drive saturation. This can lead to an increase in the reactive power absorbed by the 61 transformer, potential voltage collapses, and the creation of significant voltage harmonics 62 into the power system [Girgis & Vedante, 2015]. Power systems are affected by harmonic 63 distortion through tripping of protective relays, such as those on voltage regulation capacitor 64 banks, which can then result in power outages [Molinski, 2002]. Ultimately, the size and 65 impact of induced currents flowing in power systems are a complex interplay between 66 processes driving geomagnetic field variations [Pulkkinen et al., 2003], the proximity of the 67

power system to the auroral electrojet latitude, the power system network configuration, and
 the underlying surface conductivity structures [e.g., Divett et al., 2017].

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High peak current pulses in transformers generate leakage flux [Albertson et al., 1992] rich 71 in harmonics, increasing winding temperatures through larger eddy current losses in the 72 windings. Heating can also occur in structural parts of a transformer, with tank wall hot spots 73 subject to temperatures as high as 175°C [Kappenman, 1996]. Temperature increases within a 74 transformer are influenced by both the magnitude of the GIC, and its duration. Large GIC 75 levels have been observed to flow in power systems as a result of inter-planetary solar wind 76 magnetic field (IMF) shock events deforming the Earth's protective geomagnetic field [Fiori 77 et al., 2014; Rodger et al., 2017]. The short duration of the high GIC levels in IMF-shock 78 cases would be expected to cause significantly smaller temperature increases in transformer 79 windings than for longer duration events, potentially limiting the damage or aging of any 80 transformers exposed to these large currents [Viljanen et al., 2001]. 81

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Two additional geomagnetic drivers of GIC have been identified: substorms and 83 geomagnetic pulsations. All three drivers involve significant changes in the ground-level 84 magnetic field. Substorms have a peak occurrence around 00 MLT (Magnetic local time) and 85 can have durations of about an hour, while geomagnetic pulsations such as ULF Pc5 waves 86 can occur towards the end of a storm period, producing extended GIC periods (several hours) 87 primarily found in the MLT morning period [Pulkkinen et al., 2003]. GIC levels associated 88 with geomagnetic pulsations tend to be smaller than those associated with substorms, and 89 90 shocks.

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However, the analysis of dissolved gases (DGA) in transformer oils, which can be generated
by the breakdown of insulator materials as a result of high operating temperatures, has

suggested that there can be accumulative damage to, or aging of, transformers through 94 repeated exposure to GICs [Albertson et al., 1992]. Regular checks of DGA from 16 95 transformers in South Africa showed increased gas levels following two large geomagnetic 96 storms in November 2003 [Gaunt & Coetzee, 2007]. Three transformers tripped during the 97 storms, and three were taken out of service several months later with high DGA levels. 98 Inspection showed there was heat damage, mostly to paper insulation, in various parts of the 99 transformers. The damage was initiated by local overheating causing low temperature thermal 100 degradation as set out in Mollmann and Pahlavanpour [1999]. 101

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The currents of a transformer driven into half cycle saturation contain harmonics of various 103 orders, some even and some odd. Power networks are typically designed to cope with odd 104 harmonics as they are produced by multiple processes. In contrast, even harmonic events are 105 rare, and are uniquely produced by GIC [Ramírez-Nińo et al., 2016]. As such, networks can 106 be stressed by low-levels of even harmonics as they are not normally present [Gish et al., 107 1995]. Boteler et al. [1989] presented measurements of increased harmonic currents due to 108 GIC in a 1200 MVA transformer on the British Columbia Hydro 500 kV system. 109 Measurements were made of the fundamental and 2<sup>nd</sup> order harmonics, indicating that during 110 the GIC anticipated odd harmonics produced were by symmetrical saturation, but some very 111 strong low order even harmonics were also generated. The tripping of an overcurrent relay in 112 the Swedish power system during the large geomagnetic storm of October 2003 was triggered 113 by GIC-induced harmonic distortion [Pulkkinen et al., 2005], with the relay being 114 particularly sensitive to the 3<sup>rd</sup> harmonic. Harmonic distortion data with a time resolution of 115 one hour showed a 1.5% peak in the zero-sequence voltage (sum of the fundamental 116 frequency, 3<sup>rd</sup>, 6<sup>th</sup> harmonic etc.) at the time of the power outage. Note that there are currently 117 very few events reported with simultaneous harmonic distortion and GIC measurements. 118

Enhanced harmonics from near-by power systems have previously been observed with a 120 very low frequency (VLF) receiver during the sudden commencement of a geomagnetic 121 storm as observed in Canada [Hayashi et al., 1978]. Enhanced GIC levels giving rise to near-122 saturation of transformer core material were postulated, although no power network current 123 measurements were available at the time. Prior to the storm only odd harmonics were 124 observed, however, following the sudden commencement intense even harmonics up to 720 125 Hz (12<sup>th</sup> harmonic) appeared as well as some enhancement of the 3<sup>rd</sup> and 9<sup>th</sup> harmonics (180 126 Hz and 520 Hz). Short-lived (10s of seconds) periods with enhanced harmonic amplitude 127 were observed over the following 5 minutes as damped magnetic field oscillations occurred. 128 The highest even harmonic excited was the 18<sup>th</sup> (1080 Hz). 129

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Clilverd et al. [2010] correlated the occurrence rate of ~1 s VLF noise bursts with GIC 131 magnitude for a geomagnetic storm in May 2005. The noise bursts observed in northern 132 Finland were observed in the 20-25 kHz range and were thought to be generated by nearby 133 electrical arcing caused by the geomagnetic storm. Corroborating wideband VLF data (0.1-30 134 kHz) were not available at the time, and the correlation with GIC magnitude was from the 135 distant location of Scotland, which merely indicated that the storm was capable of creating 136 GIC. Although power line harmonic radiation has been observed extensively by VLF 137 receivers on the ground and in space [Yearby, et al, 1983; Nemec et al., 2006], very few VLF 138 wideband spectral observations of GIC transformer saturation effects associated with 139 individual space weather events have been published. 140

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In this study we present observations from the Halfway Bush substation in Dunedin, South Island, New Zealand, during two days of intense geomagnetic storm activity in September 2017. We combine and interpret a series of measurements made with high time resolution, including: GIC measurements made within the substation [Mac Manus et al., 2017],

magnetometer measurements made very close to Dunedin, broadband VLF observations 146 made at the substation, and harmonic distortion measurements made on a 110 kV high 147 voltage line linked to Halfway Bush single phase bank transformer T4. We investigate the 148 impact of short lived sudden commencement storm effects, along with longer lasting 149 substorm-like events that produced more significant responses within the substation. We 150 identify GIC levels that produced harmonic distortion, identify the components present, and 151 describe the geomagnetic perturbation characteristics. We suggest this is an unusually 152 complete set of space weather observations, providing direct evidence of GIC-saturation to a 153 transformer, along with measurements of the GIC and resultant harmonics. 154

# 155 2. Experimental Datasets

The 04-11 September 2017 interval was one of the most flare-productive periods of solar 156 cycle 24. Two active regions produced more than 15 M-class flares, and 3 X-class. A coronal 157 mass ejection (CME) associated with a 06 September X9 flare produced severe geomagnetic 158 storming on 07 and 08 September. Two solar wind shock events impacted the 159 magnetosphere, causing rapid fluctuations of the geomagnetic field. Shocks were identified 160 by SOHO at 23:13 UT on 06 September 2017 and 22:38 UT on 07 September 2017 161 [http://umtof.umd.edu/pm/]. With solar wind speeds of 600-700 km s<sup>-1</sup> the propagation time 162 for the shocks to reach the magnetosphere is in the order of  $\sim 30$  minutes. Figure 1 163 summarizes the solar wind and geomagnetic conditions for 06 to 10 September 2017 using 164 the DSCOVR measurements made at L1, and ground-based geomagnetic activity 165 measurements. An indication of the levels of geomagnetic substorm activity is given by the 166 Wp index [Nosé et al., 2012], shown in the lower panel of this figure. 167

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The arrival of the first shock late on 06 September can be seen in Figure 1 as a sudden increase in all parameters. This includes IMF  $B_z$ , suggesting that the initial disturbance is a

sudden impulse rather than a sudden storm commencement, and would therefore be only 171 weakly geo-effective as a result [Joselyn & Tsurutani, 1990]. During the first few hours of 07 172 September a recovery ensues. The second shock is seen at ~23 UT on 07 September, with 173 solar wind speeds exceeding 700 km s<sup>-1</sup> and a very strongly negative IMF B<sub>z</sub>. Such shocks 174 are known as sudden storm commencements (SSCs) and are usually geo-effective [Joselyn & 175 Tsurutani, 1990]. Finally, at ~12 UT on 08 September another strongly negative  $B_z$  period is 176 seen, and the solar wind speed remains high (>700 km s<sup>-1</sup>). A steady recovery occurs in all 177 parameters throughout the remainder of 08 September. 178

#### 179 2.1 New Zealand GIC Observations

Measurements of transformer DC neutral current have been made at the Halfway Bush 180 substation in Dunedin by Transpower New Zealand Limited since an expansion of its 181 monitoring network in the South Island in 2013 [Mac Manus et al., 2017]. The substation 182 includes the 50 Hz single phase bank transformer T4, which was notably damaged by GIC 183 effects on 6 November 2001 [Béland and Small, 2004, Marshall et al., 2012]. Neutral 184 currents were monitored using Hall-effect current transducers (Liaisons Electroniques-185 Mécaniques (LEM) model LT 505-S) with typical sampling during a GIC event of 4 s. A 186 detailed description of this dataset, along with the corrections to remove stray earth return 187 currents, was presented by Mac Manus et al. [2017]. The Halfway Bush T4 GIC observations 188 reported in this study use the corrected dataset. 189

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Total Harmonic Distortion (THD) measurements were made at circuit breaker CB532, which is connected to Halfway Bush T4 and T6 on the 110 kV bus. The percentage of THD for odd and even harmonics on each of the three phases were sampled every 10 minutes.

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#### 195 2.2 Magnetometers

The proximity of the auroral electrojet is a key factor in the characteristics of the induced 196 currents experienced by power systems [Pulkkinen et al., 2003]. Figure 2 shows an image 197 from a Suomi National Polar-orbiting Partnership satellite overpass of New Zealand from 198 13:36:13 to 13:53:32 UT on 08 September 2017. The map indicates the location of Dunedin 199 in New Zealand, and also shows bright auroral features just south of the South Island. The 200position of the bright auroral features at  $\sim 60^{\circ}$ S geomagnetic latitude (L=4) in the figure are 201 indicative of the location of the auroral electrojet [Wallis, et al., 1976] where its proximity to 202 Dunedin would be anticipated as part of a severe geomagnetic storm period which displaces 203 the electrojet equatorwards [Oughton et al., 2017]. Figure 2 also shows the location of the 204 Eyrewell (EYR) magnetometer operated by GNS Science, New Zealand which is part of 205 INTERMAGNET (http://www.intermagnet.org/). Evrewell is ~300 km from Dunedin. A 206 detailed description of the construction of EYR 1 minute averages of the horizontal 207 component of the magnetic field is given in Mac Manus et al. [2017], and the data are 208 included in this study to relate the magnetometer measurements for the current time period to 209 previous studies in this region (e.g., Mac Manus et al. 2017, Rodger et al., 2017). 210

In order to compare the Halfway Bush T4 GIC observations to changing magnetic field 211 properties associated with the nearby auroral electrojet during the severe geomagnetic storm 212 of 07/08 September 2017 we use local magnetic field measurements made at Swampy 213 Summit close to Dunedin. The locations of the Halfway Bush substation (yellow star), and 214 the Swampy Summit (University of Otago) measurement site (blue square) are shown in 215 Figure 3, plotted on a map of the local Dunedin region. The magnetometer at Swampy 216 Summit is a Bartington three-axis magnetic field sensor (MAG-03MSESL70). There is a 217 7 km separation between the Halfway Bush substation and the magnetometer. One minute 218 average magnetic field component values were calculated in a similar way to those at 219 Eyrewell. 220

Also shown in Figure 3 is the location of a magnetometer at Middlemarch (purple square), 40 km from the Halfway Bush substation. The three-axial fluxgate-type magnetometer is part to the CRUX array (http://www1.osakac.ac.jp/crux/) operated by Osaka Electro-Communication University, Japan. It was originally installed in March 2011 with 1 s sampling. The system has been described in Obana et al. [2015].

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## 227 2.3 VLF observations

Broadband VLF measurements over the frequency range 0-48 kHz were made at the 228 Halfway Bush substation site using two orthogonal vertical magnetic field loop antennas 229 connected to a computer via a sound card as described in Clilverd et al. [2009]. Although 230 there were strong local signals from the substation, the gain of the system was set to not 231 overload the aerial preamplifier or the sound card input. The time resolution of the logged 232 data was 10 s, with 46 Hz bins (48 kHz sampling using a 1024 point FFT). A similar 233 broadband VLF system was also operating at Swampy Summit 7 km away to the north, very 234 close to the site of the Swampy magnetometer. At the Swampy Summit location spectra with 235 0-48 kHz range, 10 s sampling, and 92 Hz resolution were being logged. The two orthogonal 236 VLF aerials at Swampy are orientated in north-south and east-west directions. 237

# **3. Magnetic field and GIC perturbations**

The rate of change of the horizontal magnetic field *H*-component (H' = dH/dt) at Eyrewell and Middlemarch for 06 to 09 September 2017 are shown in the upper two panels of Figure 4 (Eyrewell on the top, and Middlemarch below). The lower two panels show the Swampy Summit magnetometer H' and Halfway Bush T4 GIC for the same period. Four time periods are marked by red lines, based on the times of four distinct peaks in the rate of change of the *H*-component of the EYR data. The times of the peaks are shown in each panel as vertical red lines, and marked 1, 2, 3 and 4 in sequence. These four lines are at the following times: (1) 246 23:48 UT 06-Sep-2017, (2) 08:56 UT 07-Sep-2017, (3) 23:02 UT 07-Sep-2017, (4) 12:50 UT
247 08-Sep-2017.

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The four numbered events in Eyrewell H' (upper-middle panel) correspond well with 249 features in either the solar wind, or geomagnetic index data, described in section 2. Events 1 250 and 3 occur at the times of the two solar wind shocks arriving at Earth, with the second one 251 generating significantly larger rates of change at Eyrewell, Middlemarch and Swampy 252 Summit than the first (~35 nT/min c.f. ~7 nT/min). Following both events, the disturbances 253 gradually decline over a period of about 10 hours (the grid size is 12 hours in the panels). 254 Event 2 is associated with a small substorm following a period of intermittent IMF  $B_z$ 255 southward, and produces GIC at T4 that are insignificant enough to not be considered further 256 in this study (i.e.,  $\sim 6$  A). Event 4 occurs shortly after the strong negative deviation of IMF Bz 257 at ~12 UT on 08 September, generating Eyrewell H' rates of change of 26 nT/min, and 258 29 nT/min at Middlemarch. 259

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The Swampy Summit magnetometer data is shown in the lower left-hand panel of Figure 4. 261 Ostensibly the Swampy H' look very similar to those from Eyrewell, particularly for events 1, 262 2, and 3. However, event 4 has a much higher H' (71 nT/min c.f. 26 nT/min), and the peak 263 also occurs slightly earlier than at that of Eyrewell (i.e., 9 min beforehand). These 264 observations suggest that event 4 just after 12 UT on 08 September 2017 has localized 265 magnetic field fluctuations closer to Dunedin than Eyrewell, which is consistent with the 266 proximity of the auroral electrojet to Dunedin as identified in Figure 2. This is also consistent 267 with the power grid modelling of Butala et al. [2017], who found that the fidelity between 268 measured GIC levels and modelling using magnetometer proxies is a function of the distance 269 between the two measurements, with shorter separation distances providing more successful 270 modelling outcomes. However, Middlemarch shows a similar peak H' to Eyrewell 271

(29 nT/min c.f. 26 nT/m), which is much lower than the Swampy Summit 71 nT/min despite
there being only 37 km between the two sites. These observations clearly indicate significant
small-scale features in the magnetic field perturbations close to Dunedin, which will be
explored in more detail in Figure 5.

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The Halfway Bush T4 GIC levels show similar characteristics to the Swampy H', with 277 comparatively small induced current values for events 1 and 2, high levels for event 3, and 278 even higher levels for event 4 (up to ~49 A). Taking into account the differences in the three 279 magnetometer sites, these results suggest that event 4 generates large localized fluctuations of 280 the magnetic field, and induces large localized currents at the substation. The magnitude of 281 the GIC is similar to those shown for a geomagnetic storm on 02 October 2013 (see Figure 7, 282 Mac Manus et al., 2017), and about half of the peak GIC value thought to have damaged the 283 Halfway Bush T4 transformer during a short-duration impulse on 06 November 2001 284 [Rodger et al., 2017]. Following event 4, GIC effects continue to influence T4 for more than 285 6 hours, gradually subsiding to levels <5 A. 286

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In Figure 5 we concentrate on the two largest GIC events – event 3 and 4 - showing 288 Eyrewell, Swampy Summit, and Halfway Bush T4 data, along with the Wp index for the 289 same periods. Here the left hand column shows observations around event 3 (22:00 UT on 07 290 September to 03:30 UT on 08 September 2017) and the right hand column shows the 291 observations around event 4 (10:00 to 15:00 UT on 08 September 2017). As in the previous 292 Figure, red lines identify the times of peak H' at Eyrewell. Event 3 is triggered by the L1 293 294 solar wind shock at 22:38 UT which then arrives at Earth at 23:02 UT. The levels of H' at Eyrewell and Swampy are similar (~35 nT/min) and occur at the same time, confirming a 295 large scale perturbation of the magnetosphere as shown by the Dst variation in Figure 1. The 296 ~34 A GIC at T4 lasts only a few minutes, although GIC resulting from further magnetic 297

perturbations continue for several hours at lower levels. From 01:00 UT to 02:00 UT T4 298 shows a longer period of enhanced GIC lasting ~45 min, exceeding >15 A for about 5 299 minutes at about 01:45 UT. This response would not be predicted through inspection of the 300 equivalent Eyrewell magnetometer H' data, but can be seen as an extended period of 301 relatively high levels of H' at Swampy Summit (>10 nT/min). This enhancement of GIC, 302 occurring about 2 to 3 hours after the arrival of the solar wind shock at 23:02 UT, is 303 consistent with the effects of a geomagnetic substorm [Pulkkinen et al., 2003] as indicated by 304 the Wp index shown in the lower panel. However, the enhancement is probably only 305 noticeable in the Swampy Summit magnetometer H' data because of Dunedin's proximity to 306 the auroral electrojet, which had been displaced equatorward as a result of the prior shock 307 arrival and the associated negative IMF B<sub>z</sub>. 308

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The right hand column of Figure 5 shows observations made in the time period around event 310 4 (12:00 to 13:00 UT 08 September 2017). This event is driven by severe geomagnetic 311 storming as indicated by strongly negative IMF B<sub>z</sub> and high solar wind speed shown in 312 Figure 1. During this event the transformer T4 experienced long lasting GIC, with >15 A 313 occurring for ~30 minutes. The timing of the peak current shows good coincidence with the 314 Swampy Summit magnetometer H', but less so with the more distant Eyrewell magnetometer 315 where the peak in H' is observed ~15 minutes later. The time difference in peak H' between 316 Eyrewell and Swampy Summit appears to stem from differences in the magnitude of short-317 term variations during the substorm period, i.e., Eyrewell shows 4 peaks in H' increasing in 318 magnitude between 12:00 and 13:00 UT, whereas Swampy Summit shows peaks in H' at the 319 320 same times but with the third of the four (at ~12:45 UT) being the largest. This variation in H'responses could be due to the equatorward movement of the electrojet during the substorm, 321 with the distance from Eyrewell decreasing during the substorm time period, and the 322 electrojet passing overhead of Dunedin close to the time of the third peak (~12:45 UT). In 323

order to test this idea Pearson correlation coefficients were calculated between GIC at HWB 324 and the magnetic field observations. The calculations were performed for event 3 and event 4 325 separately, using 1 minute averaged values. The statistical significance of the Pearson 326 correlation coefficient was examined via the p-value [Martin, 2012]. For small p-values 327 (p<0.05) the null hypothesis may be rejected and the correlation assumed to be statistically 328 significant. For event 3 data were taken from 22:00 UT to 00:30 UT, 7/8 September 2017. 329 The Pearson correlation coefficient for HWB-Eyrewell was r=0.59 ( $p=10^{-15}$ ), while for 330 HWB-Swampy Summit r=0.60 (p= $10^{-16}$ ). These near-identical results indicate a large-scale 331 perturbation to the magnetic field, which generated similar effects over 100's of km. For 332 event 4 data were taken from 11:00 UT to 03:30 UT, 8 September 2017. The Pearson 333 correlation coefficient for HWB-Evrewell was r=0.38 ( $p=10^{-11}$ ), while for HWB-Swampy 334 Summit r=0.55 (p= $10^{-25}$ ). These results are consistent with small-scale structures being 335 present and a decrease in correlation with increasing distance from HWB. In all of these cases 336 the p-values are very close to zero, indicating statistically significant correlations. The Fisher 337 Z-test [Preacher, 2002] was also undertaken in order to test whether the differences between 338 the correlation coefficients found using data from the two different magnetometer sites in 339 event 3 are significant, and similarly for event 4. The difference between two independent 340 correlation coefficients are considered significant if z > modulus(1.96). For event 3, z=-0.133341 (for 151 data points), i.e. there is no significant difference between the two correlation 342 coefficients (0.59 and 0.6). While for the event 4, z=-2.66 (for 301 data points), i.e., the 343 difference between the correlation coefficients (0.38 and 0.55) is significant. This test 344 confirms that the decrease in correlation with increasing distance from HWB for event 4 is 345 346 significant.

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Unlike the GIC generated by the arrival at Earth of the two solar wind shocks, event 4 shows a gradual enhancement in GIC levels over about 15 minutes. This enhancement is

probably caused by expanding localized ionospheric current systems at ~110 km associated with substorm activity [Pulkkinen et al., 2003] and is co-incident with the substorm event that began just after 12 UT as shown by the Wp index in the lower panel. The event timing of 12-13 UT (01-02 MLT) is consistent with substorm occurrence close to local midnight MLT. This event is the most extreme and long lasting of the GIC recorded at the T4 transformer during the 07/08 September 2017 geomagnetic disturbances.

# 356 **4. Harmonic distortion observations**

The broadband VLF magnetic field observations undertaken at the Halfway Bush substation 357 showed two periods of enhanced harmonic intensity during the geomagnetic disturbances. 358 The same periods were also detected by a similar broadband VLF system operating at 359 Swampy Summit 7 km away. The observations are shown in Figure 6. The two upper panels 360 show the Halfway Bush wideband frequency-time data taken around event 3 on the left 361 (22:00 to 03:30 UT on 07/08 September), and around event 4 on the right (10:00 to 15:00 UT 362 on 08 September). The frequency range is 0-5 kHz, with a frequency resolution of 46 Hz, and 363 10 s sampling. The color scale represents signal intensity relative to an arbitrary level. As in 364 our previous figures the red vertical lines indicate the times of peak H' at Eyrewell. The 365 middle panels show the same data from Swampy Summit, but at 92 Hz resolution. Many 366 horizontal lines can be seen in all panels. These represent harmonics generated by localized 367 loads on the substation, and include 3 intense pulses at  $\sim 1$  kHz which are most likely due to a 368 1050 Hz ripple injection system used to control domestic hot water heating systems in the 369 Dunedin area (01:30, 11:00, 11:15 UT on 08 September). Ripple control in Dunedin was 370 previously described by Werner et al. [2005]. However, at 01:45 UT and 12:30 UT 371 enhancements of harmonics up to 1.6 kHz are seen at Halfway Bush, and up to 1 kHz at 372 Swampy Summit. These times are co-incident with high GIC levels shown in the lower 373 panels. At Swampy Summit the enhanced harmonics were detected in the directional north-374

south aerial, while nothing was seen in the equivalent east-west aerial. As Swampy Summit is
located almost directly north of the Halfway Bush substation this observation is regarded as
confirmation that the harmonics were detected directly from the substation rather than via
transmission through nearby power lines.

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The enhanced harmonics observed are closely associated with time periods when transformer T4 experienced greater than ~15 A GIC. The duration of the two periods also match the length of time that the GIC exceeds 15 A, i.e., 30 minutes during event 4, and 5 minutes during the substorm that followed event 3. Notably event 3 itself, which generated ~34 A at 23:00 UT on 07 September shows no evidence of any enhanced harmonics, which is presumably due to the short-lived nature of the shock-induced magnetic field perturbations.

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A more detailed analysis of the 08 September 2017 wideband data is shown in Figure 7. The 387 upper panel shows the average amplitude of each 46 Hz frequency bin in the range 0-2500 Hz 388 for the 30 min period prior to the harmonics being present in event 4 (11:45-12:15 UT) 389 compared with the 30 min period where harmonics were present (12:15-12:45 UT). The panel 390 shows that during event 4 the enhanced harmonics spanned 50-1600 Hz. The middle panel 391 shows the difference between the two lines presented in the upper panel. Six harmonics that 392 were significantly enhanced during the GIC period are indicated by red vertical dashed lines, 393 i.e., the 300 Hz (6<sup>th</sup>), 600 Hz (12<sup>th</sup>), 950 Hz (19<sup>th</sup>), 1200 Hz (24<sup>th</sup>), 1350 Hz (27<sup>th</sup>), and 394 1500 Hz (30<sup>th</sup>) harmonics. Some harmonics were already present prior to event 4, however, 395 there is little evidence of even harmonics in the VLF spectra, as expected. The harmonics 396 already present included the 750 Hz (15<sup>th</sup>) harmonic, which showed no change at the time of 397 the enhanced GIC. However, harmonics in the range 950-1250 Hz were already present and 398 were enhanced during the GIC period. These observations are consistent with previous work 399 [Hayashi et al., 1978] showing the appearance of strong even harmonics (in our case numbers 400

6, 12, 24, and 30), and the enhancement of some already existing odd harmonics (numbers 19
and 27 in this example) during GIC events.

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The lower panel of Figure 7 shows the VLF signal intensity from 0 to 2000 Hz, averaged 404 with 1 minute resolution. The panel shows an increase in intensity above typical background 405 levels, from 12:15 to 12:45 UT, consistent with the T4 GIC levels exceeding 15 A and non-406 linear saturation beginning to occur. The temporal variation in 0-2000 Hz average intensity is 407 also consistent with the variation in T4 GIC levels shown in Figure 5, notably the two peaks 408 that occur at ~12:30 and ~12:45 UT. This suggests a close correlation between the levels of 409 GIC in the substation and the <2 kHz VLF intensity, with the VLF observations also 410 providing harmonic information at the same time. 411

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The strong north-south directionality of the harmonics detected at Swampy Summit, 7 km 413 due north of Halfway Bush, suggest that the signals originate from the local distribution 414 network. Harmonic currents can flow in either the transmission system or the local 415 distribution network. Transmission line relative impedance defines which will happen, 416 particularly in the case of resonant frequencies which can experience high impedance for 417 certain line-lengths. As a result the harmonic current is forced into the on-site distribution 418 network [Carroll et al., 1993]. The VLF system appears to have the potential to determine 419 which harmonics are affected in this way, and the large increase in the amplitude of the 420 1350 Hz (27<sup>th</sup>) harmonic suggests that it is a local resonant frequency. 421

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It is possible to relate the VLF observations of enhanced harmonics from the Halfway Bush substation with total harmonic distortion (THD) measurements of the 110 kV bus connected to Halfway Bush T4, made at circuit breaker CB532 connected to Palmerston. This is depicted in the Halfway Bush single line diagram shown in the upper panel of Figure 8,

which shows the high voltage connections at the substation. The lower panels of Figure 8 427 shows the odd THD (left panel) and the even THD (right panel) for the disturbance period 428 from 12:00 UT, 06 September 2017 to 00:00 UT, 10 September 2017. As before, red vertical 429 lines indicate the event times of peak H' at Eyrewell. The three components of the 3-phase 430 electrical power system THD voltage are shown, following the standard naming convention 431 of red, yellow, and blue phases, each alternating current carried by a separate conductor but 432 offset from each other by one third of a cycle, (here denoted A, B, and C respectively). The 433 odd harmonics show a diurnal variation around 1% distortion, with smaller levels of 434 variability occurring particularly during the daytime in Dunedin (00 UT  $\pm$  6 hrs). Small 435 percentage perturbations of the order of 0.1-0.2% occur for the first, third and fourth GIC 436 events. Even harmonic THD is rarely above a level of 0.05%, but does show peaks that 437 coincide with the third event, the substorm following 3 hours after the third event ( about 438 01:45 UT, 08 September), and very obviously for the fourth event. The largest even THD 439 recorded was  $\sim 0.9$  % while the odd THD at the same time was  $\sim 0.8$ %, which should be 440 compared against the steady state upper THD recommendation of 2.5% for 69-161 kV [IEEE 441 Std 519-2014]. 442

443

Detailed even and odd THD changes during events 3 and 4 (with the same time ranges as in 444 Figures 4 and 5) are shown in Figure 9. The upper panels show the odd harmonic distortion, 445 while the lower panels show the even harmonic distortion percentages. The red vertical lines 446 indicate the same timing as in previous figures. Event 3 at 23:00 UT on 07 September shows 447 little clear association with any distortion, odd or even, while the substorm at 01:45 UT leads 448 449 to a decrease in odd distortion of ~0.1 percent at the same time as an increase to 0.4% in even THD. A similar decrease/increase response can be seen for event 4 at 12:30 UT. Odd THD 450 decreased by ~0.3% while the even THD increased to as much as ~0.9%. The duration of the 451 elevated even THD during event 4 is 30-40 minutes (with 10 minute sampling) over the time 452

period 12:20 UT to 12:50 UT, which agrees well with the timing of the VLF broadband
observations of harmonics, shown earlier.

455

The observations during this period demonstrate evidence of transformer stress (through 456 generated harmonics) linked to the magnitude of GIC due to changes in the local magnetic 457 field. The data presented in this study is consistent with existing indirect evidence that 458 geomagnetic storms impact electrical network performance in New Zealand. A study 459 examining ~4 years of New Zealand real time wholesale electricity prices and geomagnetic 460 disturbances concluded that the "standard deviation in the nodal price, a measure of energy 461 losses and transmission constraints, is more likely to be large when the GIC proxy is large" 462 [Forbes & St. Cyr, 2008]. The study reported that real-time market conditions for power grids 463 in multiple locations around the world were statistically related to geomagnetic disturbances 464 (a GIC proxy), and that this effect was strongly seen in the New Zealand data. 465

466

## 467 **5. Summary and Discussion**

In this study we have analyzed the effects of a geomagnetically disturbed period on the Halfway Bush substation located in Dunedin, New Zealand. There was unprecedented data coverage from a unique set of instruments, i.e., measurements of GIC effects on the T4 transformer housed within HWB, magnetic field measurements from a magnetometer located only 7 km from HWB, VLF wideband measurements at, and near, HWB detecting the presence of power system harmonics, and voltage distortion measurements made on a 110 kV line circuit breaker connected to the T4 transformer.

475

During 06 and 07 September 2017 two interplanetary solar wind shocks occurred, generating four distinct periods of GIC. The two shock events produced very different

levels of disturbance, probably due to the orientation of the vertical component of the IMF, *Bz*. Both shocks were closely associated in time with short lived spikes in GIC, and were followed by declining effects over 6-10 hours consistent with the effects of geomagnetic pulsations, but neither shock generated any observable harmonics at HWB, indicating that the substation transformers were not significantly stressed by the events, even though T4 reported a short-duration peak GIC of 34 A during the second shock event.

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However, two periods of substorm activity generated longer-lasting GIC in T4, with 485 harmonic distortion detected from the substation when GIC levels exceeded about 15 A. 486 The observations indicated that half-wave saturation of the transformer occurred 487 continuously for 30 minutes during the most effective substorm. VLF receiver systems 488 picked up harmonics up to the 30<sup>th</sup> harmonic, with only specific harmonics enhanced 489 possibly due to in-line impedance resonances in the transmission system. The VLF 490 harmonic observations were consistent with increases in even harmonic THD detected on a 491 connecting 110 kV line, which occurred co-incidentally with small decreases in odd 492 harmonic THD. 493

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The substorm periods showed better agreement between T4 GIC variations and nearby Swampy Summit magnetometer data, than for the more distant Middlemarch and Eyrewell magnetometer data. This suggests that small scale ionospheric currents were important in the substation response to geomagnetic activity levels. However, for the global scale solar wind shock events, data from all three magnetometer sites (Swampy Summit, Middlemarch, and Eyrewell) showed similar temporal behaviour to the relative variations of the HWB T4 GIC.

502

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https://zenodo.org/record/1246869#.WvoINXWFOK5. DSCOVR data can be found at
https://www.ngdc.noaa.gov/dscovr/portal/index.html#/. Dst and Wp data are available at
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**Figure 1.** A summary plot of the solar wind and geomagnetic conditions during the disturbed period in September 2017. DSCOVR solar wind speed Vx, and IMF magnetic field components (Btot and Bz) are shown in the upper two panels. Geomagnetic activity index Dst, and substorm activity index Wp are shown in the lower two panels. Vertical dotted blue lines indicate the times of interplanetary shocks reported by the SOHO spacecraft.



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**Figure 2.** Image provided by a Suomi National Polar-orbiting Partnership satellite overpass of New Zealand, very shortly after the period of peak GIC on 08 September 2017 in the local time early morning. The image was generated by the Near Constant Contrast product made by the Day Night Band of the VIIRS sensor. The VIIRS image granules shown here span from 13:36:13 to 13:53:32 UT, with the spacecraft moving north to south. The locations of Eyrewell and Dunedin, South Island, are indicated by squares. Contours of constant geomagnetic latitude are shown as solid, dot-dashed, and dotted lines.

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Figure 3. Map of the Dunedin region showing the relative locations of the Halfway Bush (HWB) substation as a yellow star, and the University of Otago research station in the hills of Dunedin at Swampy Summit (blue square). In addition, the Osaka University magnetometer at Middlemarch ~37 km from Swampy is also plotted (magenta square). Note the overall location of Dunedin is also shown in Figure 2.



**Figure 4.** Magnetic field observations from Eyrewell (upper panel), Middlemarch and Swampy Summit (middle panels), and Halfway Bush transformer T4 GIC magnitude measurements (lower panel) across the 06-09 September 2017 geomagnetically disturbed period. Four time periods are marked by red lines, based on the times of four distinct peaks in the rate of change of the H-component of the EYR data. These times are shown in each panel.



Figure 5. Detailed examination of the magnetic field rate of change and GIC variations during the large events 3 (left-hand column) and 4 (right-hand column) identified in Figure 4. Upper panels: Eyrewell H'; middle panels: Swampy Summit H' and Halfway Bush T4 GIC magnitude; lower panel: Wp index. Red lines indicate the times of peaks in H' at Eyrewell.





**Figure 6.** Wideband VLF observations from beside the HWB substation and at Swampy Summit north-south aerial. Event 3 is shown in the left column and event 4 is shown in the right column. Periods at about 01:45 UT, and 12:30 UT show evidence of harmonic distortion produced during the high current, long lasting GIC events. HWB GIC data are plotted in the lower panel in order to provide easy comparison with the VLF data. Vertical red lines indicate times of large H' at Eyrewell.



Figure 7. Upper panel. Wideband VLF observations from the HWB substation on 08 732 September 2017, showing the mean amplitude for each frequency bin before (11:45-12:15 733 UT, blue line) and during (12:15-12:45 UT, red line) the GIC-induced enhanced harmonics 734 period. Increases in harmonics can be seen up to ~1600 Hz. Middle panel. The difference 735 in received amplitude before and during the GIC>15 A period. Substantial increases are 736 seen at several frequencies, notably 300, 600, 950, 1200, 1350, and 1500 Hz, while some 737 harmonic frequencies show little change (i.e., 750 and 1250 Hz). Lower panel. The average 738 intensity of the wideband signal in the range 0-2000 Hz, showing an increase from 12:15 to 739 12:45 UT. 740



Figure 8. Upper panel. A single line diagram of the Halfway Bush substation high voltage
connections, including the circuit breaker CB532 on a 110 kV connection to Palmserston.
Lower panels: Total Harmonic Distortion (THD) measurements made at HWB on a 110 kV
line across the 06-09 September 2017 geomagnetically disturbed period. The percentage of
THD for odd harmonics (left panel) and even harmonics (right panel) are shown separately.
The same four time periods defined from Figure 4 are also indicated. Data representing the
red, yellow, and blue phases (denoted A, B, and C respectively) are shown.



**Figure 9.** Total Harmonic Distortion (THD) measurements made at HWB for the voltage on a 110kV line during the same time periods shown in Figure 4 and 5. The average THD percentage for odd harmonics (upper panels) and even harmonics (lower panels) are shown separately. Data representing the red, yellow, and blue phases (denoted A, B, and C respectively) are shown. Vertical red lines indicate times of large *H*' at Eyrewell.