

Investigating radiation belt losses through numerical modelling of precipitating fluxes

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Abstract. It has been suggested that whistler-induced electron precipitation (WEP) may be the most significant inner radiation belt loss process for some electron energy ranges. One area of uncertainty lies in identifying a typical estimate of the precipitating fluxes from the examples given in the literature to date. Here we aim to solve this difficulty through modelling satellite and ground-based observations of onset and decay of the precipitation and its effects in the ionosphere by examining WEP-produced Trimpi perturbations in subionospheric VLF transmissions. In this study we find that typical Trimpi are well described by the effects of WEP spectra derived from the AE-5 inner radiation belt model for typical precipitating energy fluxes. This confirms the validity of the radiation belt lifetimes determined in previous studies using these flux parameters. We find that the large variation in observed Trimpi perturbation size occurring over time scales of minutes to hours is primarily due to differing precipitation flux levels rather than changing WEP spectra. Finally, we show that high-time resolution measurements during the onset of Trimpi perturbations should provide a useful signature for discriminating WEP Trimpi from non-WEP Trimpi, due to the pulsed nature of the WEP arrival.

Key words. Ionosphere (ionization mechanisms) – Magnetospheric physics (energetic particles, precipitating; magnetosphere-ionosphere interactions)

1 Introduction

It has been suggested that whistler-induced electron precipitation (WEP) may be the most significant inner radiation belt loss process for some electron energy ranges (e.g. Dungey, 1963; Rodger et al., 2003). One area of uncertainty lies in providing a typical description of the precipitating fluxes. Here we aim to solve this difficulty through modelling satellite and ground-based observations of onset

and decay of the precipitation and its effects in the lower ionosphere. Whistler-induced electron precipitation from the Van Allen radiation belts occurs as a result of coupling between the troposphere and the magnetosphere. The energetic electron precipitation arises from lightning produced whistlers (Storey, 1953) interacting with cyclotron resonant radiation belt electrons near the equatorial zone (Tsurutani and Lakhina, 1997). Pitch angle scattering of energetic radiation belt electrons (Kennel and Petschek, 1966) by whistler mode waves drives some resonant electrons into the bounce loss cone, resulting in their precipitation into the atmosphere (Rycroft, 1973). An important parameter for determining the overall importance of WEP to radiation belt losses is the magnitude of a “typical” WEP event. This may be calculated from theoretical studies (e.g. Abel and Thorne, 1998) or inferred from experimental observations, such as in-situ measurements of WEP events (Voss et al., 1998). Estimates of electron lifetimes due to WEP-driven losses using ground-based whistler observations as a proxy for WEP events took a typical WEP event as removing 0.0004% of the trapped flux (Burgess and Inan, 1993), and found that these losses might be comparable with those from plasmaspheric hiss. However, a similar study estimated that typical WEP events might remove only 0.00001% of the trapped flux (Smith et al., 2001), and hence that whistlers were not significant in overall inner-belt losses. A different approach has been to use experimental observations to characterize typical WEP magnitudes. Combining reports of satellite WEP observations with ground-based whistler measurements, Rodger et al. (2003) concluded that the typical WEP mean precipitation energy flux was 2×10^{-3} ergs $\text{cm}^{-2} \text{s}^{-1}$, and that in some electron energy ranges WEP may be the most significant inner radiation belt loss mechanism.

One complementary technique to study WEP makes use of long-range remote sensing of very low frequency (VLF) waves propagating inside the waveguide bounded by the lower ionosphere and the Earth’s surface. Significant variations in the received amplitude and/or phase of fixed frequency VLF transmissions arise from localized changes in

the lower ionosphere. Further discussion on the use of subionospheric VLF propagation as a remote sensing probe can be found in recent review articles (e.g. Barr et al., 2000; Rodger, 2003) and studies (Cummer and Inan, 2000; McRae and Thomson, 2000; Bainbridge and Inan, 2003). WEP leads to localized ionospheric modifications produced by secondary ionisation just below the D-region of the ionosphere, which are observed as “Trimpi” perturbations in subionospheric VLF transmissions (Helliwell et al., 1973). These perturbations begin with a relatively fast (~ 1 s) change in the received amplitude and/or phase, followed by a slower relaxation (< 100 s) back to the unperturbed signal level due to the recombination of the additional ionisation. Trimpi perturbations permit observers to study WEP fluxes and the chemistry of the nighttime lower ionosphere (e.g. Pasko and Inan, 1994), from locations remote from the actual precipitation region.

Until recently, there was considerable uncertainty as to the typical size of the D-region patch altered by WEP. Trimpi perturbations have been used to show that WEP produced patches are large (at least $600 \text{ km} \times 1500 \text{ km}$) (Clilverd et al., 2002), considerably larger than the observed dimensions of whistler ducts (e.g. Angerami, 1970). Large D-region patch dimensions have been explained through a quasi-trapped whistler propagation theory in which ducted energy spreads at the magnetic equator (Strangeways, 1999), resulting in whistler-mode signals which have leaked outside their whistler duct still contributing to the horizontal lateral extent of WEP. Such leakage would give a significantly larger precipitation footprint than the actual dimensions of the whistler duct. A different mechanism also leading to large WEP patch dimensions comes through the precipitation caused by obliquely (nonducted) propagating whistlers, creating an ionospheric disturbance of $\sim 1000 \text{ km}$ spatial extent (Johnson et al., 1999). At this stage there is no clear experimental evidence to indicate whether quasi-ducted or nonducted whistler propagation dominates the overall WEP losses (Rodger et al., 2003).

The decay time scale information from Trimpi signatures has been used to determine the vertical dimensions of the WEP modified ionospheric patches. An analysis of 134 Trimpis observed at Palmer station (Antarctica) on transmissions from NPM (Hawaii) reported that the “recovery signatures” (decay) could be approximated by an exponential function for the purposes of determining the decay rate (Pasko and Inan, 1994). A later report concluded that the decay of Trimpis follow a logarithmic dependence (Dowden et al., 2001), noting that the difference between exponential decay and logarithmic decay is relatively small, except near the beginning (near onset) and end of the perturbation. This study also noted that a re-examination of Trimpi observed in Tokyo from NWC (Australia), were also consistent with a logarithmic decay signature, rather than an exponential decay with two characteristic time scales, as originally concluded (Molchanov et al., 1998).

A number of theoretical studies have made use of electron density perturbations with a vertical Gaussian profile

to represent the WEP produced ionospheric electron density modification (Nunn and Strangeways, 2000; Clilverd et al., 2002), while others have studied modifications derived from satellite observed WEP fluxes (e.g. Pasko and Inan, 1994; Rodger et al., 2002). For example, it has been reported that the observed Trimpi recovery signatures may be used to determine the energy content of WEP bursts (Pasko and Inan, 1994). In this paper we closely examine the onset and decay of Trimpi perturbations so as to better understand the WEP fluxes producing the ionospheric modification, characterized through experimental WEP observations, and thus investigate the estimates of the likely impact of the WEP on the radiation belts.

2 Modeling Trimpis due to WEP impact

In this section we define the process through which we attempt to understand the nature of precipitation fluxes by modelling a specific set of precipitation conditions (transmitter-receiver great circle path) described by Clilverd et al. (2002), combining flux spectrum observations, a model of the atmosphere, and a VLF scattering model.

2.1 Scattering model and situation

We consider the VLF Trimpi perturbations that have been observed at the Faraday research station (65.3° S , 64.3° W , $L=2.5$), Antarctica, on transmissions from NPM (23.4 kHz, Oahu, Hawaii). The WEP is taken to produce an ellipse-shaped ionospheric modification with dimensions $850 \times 2150 \text{ km}$, centred 800 km away from Faraday on the 12.3 Mm NPM-Faraday great circle path (GCP), the ellipse having its long axis orientated in the magnetic E/W direction as shown in Fig. 1. The WEP-modified ionospheric modification is termed a Lightning Induced Enhancement (LIE) “patch”. This position and ionospheric modification size are taken from Clilverd et al. (2002), and are determined by observations at multiple Antarctic Peninsula stations. Inside the WEP modified ionospheric region the perturbation to the collision frequency profile is assumed to be zero due to rapid plasma cooling ($\sim 0.1 \text{ s}$ at 85 km , and faster for lower altitudes; Rodger et al., 1998), and thus the modification is described solely in terms of perturbations to the electron number density profile $dN_e(x, y, z)$, determined in Sect. 2.2 and taken to be independent of horizontal coordinates x and y inside the LIE ellipse.

The transmitter is modelled as a vertical electric dipole at zero altitude. The receiver at Faraday is assumed to be a vertical electric dipole antenna, though in experimental practice the horizontal magnetic fields are measured with a loop antenna. Apart from the WEP modified ionospheric patch the background ionosphere is assumed to be homogeneous along the GCP. The unperturbed ionospheric electron density and collision frequency profiles are given by a standard nighttime

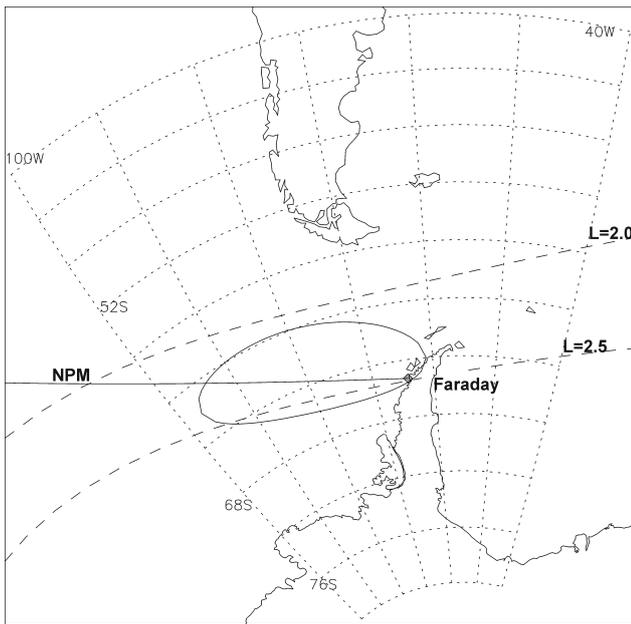


Fig. 1. Map of the region around the Antarctic Peninsula, showing the location of Faraday Station. The Great Circle Path to the transmitter NPM is plotted, along with the $L=2.0$ and $L=2.5$ contours at 100 km altitude. An ellipse marks the WEP modified ionospheric region used in our calculations.

ionospheric model. The effective collision frequency (ν) profile is given as a function of vertical coordinate z , by

$$\nu(z) = 1.816 \times 10^{11} e^{-0.15z} \quad [\text{s}^{-1}], \quad (1)$$

while the unperturbed electron number density profile, $N_e(z)$, at D-region altitudes is specified by a Wait ionosphere (Wait and Spies, 1964),

$$N_e(z) = 7.855 \times 10^{-5} e^{\beta(z-h')} \nu(z) \quad [\text{el. cm}^{-3}]. \quad (2)$$

Here β represents the “sharpness” of the lower ionospheric boundary and h' represents the effective height of the ionosphere, taken as being $h'=87$ km and $\beta=0.5$ km $^{-1}$ (dashed line) for typical nighttime conditions. The neutral atmosphere is described using the MSIS E-90 model atmosphere (Hedin, 1991) for (62.5° S, 280° E, $L=2.23$), roughly the centre of the patch, for 06:00 UT, 23 April 1994. The ground’s electrical properties also are assumed to be homogeneous and to have the properties of seawater, with conductivity 4 S m $^{-1}$, and dielectric constant of 81.

2.2 Production of ionospheric modification from WEP

The differential number spectrum as a function of energy of our “typical” WEP burst, $N(E)$, is found as described in Rodger et al. (2003), assuming an electron beam limited to the range 1–1500 keV, produced by a whistler spanning 0.5–5 kHz with power spectral density as given by the lightning spectra. It is assumed that the WEP lasts 0.2 s, and has a mean precipitation energy flux of 2×10^{-3} ergs cm $^{-2}$ s $^{-1}$. A cold plasma density from Menk et al. (1999) is taken. In our

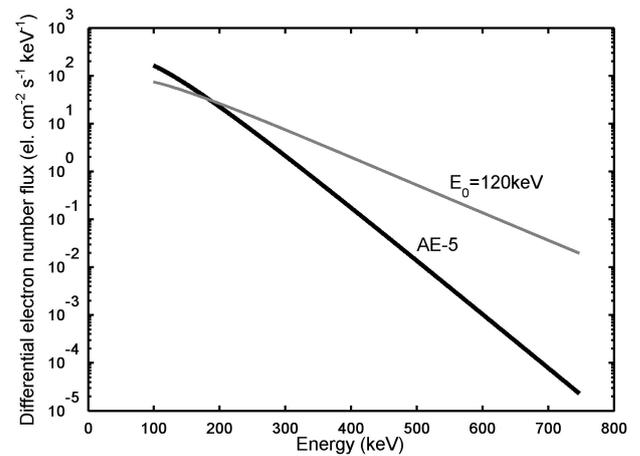


Fig. 2. Differential electron number flux of precipitated electrons in WEP events versus energy, showing the two different energy spectra used to characterize the bursts. The heavy line uses the empirical AE-5 trapped electron model, while the lighter grey line uses an e -folding energy of 120 keV.

initial calculations we will consider WEP bursts with two different energy spectra, as shown in Fig. 2. In the first case A the WEP energy spectrum is determined from the spectrum of the trapped electrons in the pitch angle range from the loss cone angle (α_{LC}) to ($\alpha_{LC} + 0.5^\circ$) given by the empirical AE-5 Inner Zone Electron Model (Teague and Vette, 1972), after Rodger et al. (2003). In the second case B the WEP energy spectrum is taken to be proportional to $\exp(-E/E_0)$, where E is the kinetic energy of a trapped electron and E_0 is the e -folding energy, with $E_0=120$ keV. This value is taken from a detailed analysis of an experimentally observed WEP event (SEEP Event (D); Voss et al., 1998), which concluded that the differential energy spectrum of the trapped electrons for this event could be best described by $E_0=120 \pm 40$ keV (Inan et al., 1989). Note that this e -folding energy is roughly two and a half times bigger than that found for case A, leading to a WEP burst (B) having a greater high-energy contribution, as seen in Fig. 2.

The ionospheric effects due to each type of WEP burst may be determined by convolving the ionisation rate for a monoenergetic electron beam with the differential number spectra, $N(E)$ shown in Fig. 2, and integrating over the energy range 1–1500 keV. The WEP burst ionisation rate $q(E)$ is found by an application of the expressions in Rees (1989), as has been previously described by Rodger et al. (2002).

2.3 Relaxation of WEP-modified regions

The relaxation of the WEP-induced ionisation is calculated from the end of the WEP burst, which is taken to be at $t=0$, assuming no changes to the WEP ionospheric modifications during the burst, i.e. that the relaxation time is significantly longer than the 0.2-s burst. Electron loss rates are calculated using reasonable values for the attachment rates and recombination coefficients at ionospheric altitudes (Rodger

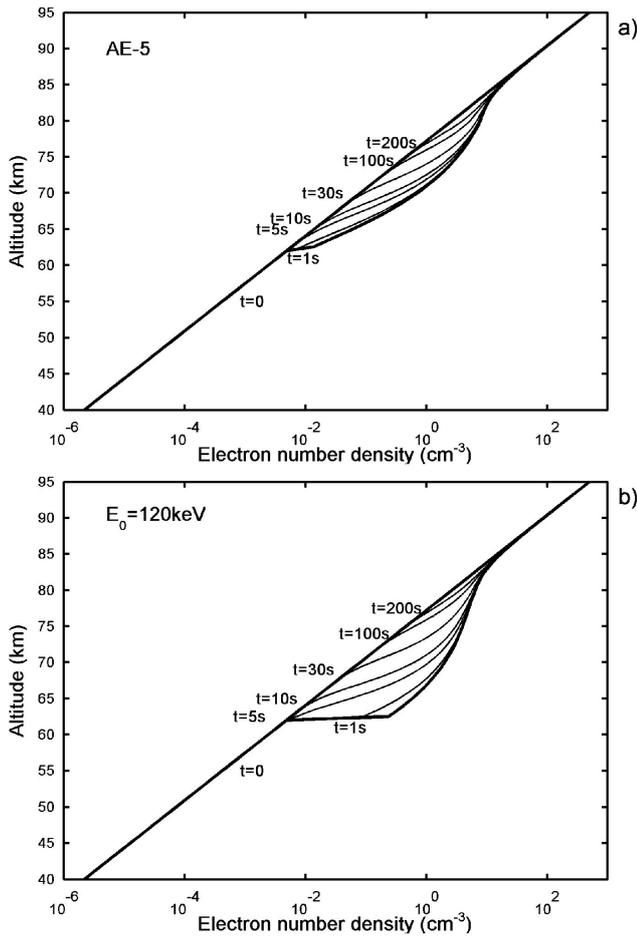


Fig. 3. Relaxation of different ionisation changes in the lower ionosphere at various times after the creation of the WEP-produced patch. The initial change at $t=0$, and the ambient night-time electron density profile, are shown as heavy lines. (a) Ionisation produced by the AE-5 WEP burst to the lower ionosphere. (b) Ionisation due to the e -folding profile.

et al., 1998), previously used to investigate red sprite associated ionisation (Nunn and Rodger, 1999) and solar eclipses (Clilverd et al., 2001). Ambient electron production rates (separate from WEP) are calculated by assuming that the ionosphere is initially in equilibrium. As such, the simulated ionospheric electron density will relax towards the ambient profile from any perturbed state. The loss rates given by the Rodger et al. (1998) expressions depend on the temperature of the electrons and neutrals present and on the instantaneous electron density. While WEP events are likely to cause significant changes in electron temperature, these will be short-lived in comparison with modifications to the electron density (Rodger et al., 1998). Thus, throughout our simulation, the electron temperature is taken to be equal to the neutral temperature.

The relaxation of the ionisation changes for the WEP-pulses (Fig. 2) are shown in Fig. 3. In both panels of this figure the heavy line shows the total electron density caused by the 0.2-s WEP burst at $t=0$, before the ionisation starts

to relax towards the ambient profile (also shown as a heavy line). A number of intermediate ionisation profiles are also shown on the figure describing the ionisation profiles for various times after the WEP disturbed patch is created. In reality, the ionisation profiles $N_e(z)+dN_e(z,t)$ have been determined with higher time resolution than shown in this figure; in the first 30 s the maximum time resolution of the ionisation profiles is 0.4 s, followed by longer, nonlinear steps at the tail-end of the perturbation.

2.4 Scattering theory and implementation

The linear Born theory of 3-D VLF scattering from an ionospheric plasma perturbation is fully described in Nunn (1997). The Born approximation involves assuming that the total incident field at a given point in the LIE is the incident zero order field (\underline{E}_0). As shown by Nunn (1997), a better linearization may be achieved by assuming that the perturbation of the electric displacement vector is zero within the LIE rather than that the perturbation of the electric field is zero. The modified region is covered by a 3-D spatial grid, and the zero-order incident field \underline{E}_0 calculated for each grid point. VLF propagation from the transmitter to each grid point is computed using modal theory and the Naval Ocean Systems Center (NOSC; San Diego, USA) MODEFNDR code (Morfit and Shellman, 1976). The MODEFNDR code returns all the parameters of the subionospheric VLF modes, as described by Wait (1996). MODEFNDR receives as inputs the frequency, as well as the ionospheric and geomagnetic parameters. Then, assuming horizontal homogeneity, but taking into account the curvature of the Earth, the program calculates the appropriate full wave reflection coefficients for the waveguide boundaries and searches for those complex modal angles which give a phase change of 2π across the guide. At each grid point, linear scattering theory furnishes an effective current $\underline{J}_{eff}(t)$ which acts as the source of the scattered field \underline{E}_s . Clearly, for large LIE patches and/or patches with large changes in electron number density the Born approximation will break down, since the total field at a point in the LIE will not be well approximated by the zero-order incident field. By comparing with a 2-D Finite Element modelling of VLF scattering, it was shown in Nunn et al. (1998) that for an LIE with horizontal dimensions ~ 100 km the Born approximation fails for maximum electron density perturbations >6 el. cm^{-3} .

The current version of the code neglects all components of the conductivity tensor, except the zz component. When the transmitter and receiver are both vertical electric dipoles, this is amply justified to an overall accuracy of the order of a few percent (Clilverd et al., 1999). The vertical electric field at the receiver is computed for each elementary volume of the modified region using MODEFNDR, summed over the total modification volume to give the total scattered field, and thus the observed VLF perturbation. The height gain functions returned by MODEFNDR become inaccurate at high altitudes, although generally for those altitudes above the VLF reflection height. Accordingly, an attenuation function has been

applied to these functions to prevent an incorrect contribution appearing due to ionisation at heights >85 km. This attenuation function is piecewise linear and was derived from a separate calculation that integrates the Booker quartic through the ionosphere for the dominant TM VLF modes. Contributions to VLF scattering from high altitudes are therefore estimates based on the attenuation function applied to the height-gain functions. We expect any inaccuracies to be small, as the presence of a significant high altitude scattering contribution would lead to a long tail in the Trimpri time signature, which is not seen in experimental observations. In future studies it is proposed to replace MODEFNDR with a full wave propagation code or a substantially modified version of MODEFNDR. This code has previously been used to study the size of WEP produced ionospheric patches (Clilverd et al., 2002), and the relaxation of ionospheric ionisation produced by red sprites (Nunn and Rodger, 1998).

2.5 LIE characteristics based on WEP bursts

Using the ionisation profiles $N_e(z)+dN_e(z,t)$, in association with the 3-D scattering code, we have calculated the time-varying NPM VLF perturbations expected at Faraday due to the relaxing ionospheric modifications shown in Fig. 3. The resulting perturbations from the undisturbed wave due to an AE-5 WEP burst (case A) are shown in the top panel of Fig. 4. The heavy black line shows the perturbation to the received amplitude, while the dashed line gives the perturbation in the received phase. As expected for perturbations due to large homogenous modifications straddling the GCP, the transmission amplitude decreases while the transmission phase increases (Wait, 1964). As the ionospheric modification relaxes, the perturbations decay back to the undisturbed level. Note that the phase perturbation returns to the unperturbed level considerably more slowly than the amplitude perturbation, as observed in experimental data (e.g. Dowden et al., 2001). This phenomenon can only be due to the phase of the scattered signal relative to the direct signal approaching an angle of $\sim 90^\circ$ as the perturbation decays away. There is nothing in the theory suggesting 90° as a limiting value, and this would appear to be somewhat coincidental. Although the start of the VLF perturbation is shown in the figure, the onset is not very meaningful, as the calculations have been undertaken directly after the arrival of the WEP burst. We calculate the effects of the perturbation onset in detail in Sect. 5.

In order to consider the decay of the Trimpri perturbation we do not focus upon the amplitude or the phase perturbation, which are often quite different, but rather the vector combination through the standard phasor diagram (e.g. Dowden and Adams, 1988). This produces the phasor of the wave scattered off the WEP burst plasma, which, added to the unperturbed phasor of the NPM transmission, gives the phasor of the perturbed wave. We call this the scatter phasor, which has magnitude, M , and phase, ϕ , (to avoid the confusing nomenclature of “scatter phasor phase”, ϕ has often been referred to as the “echo phase”). As seen in the middle panel

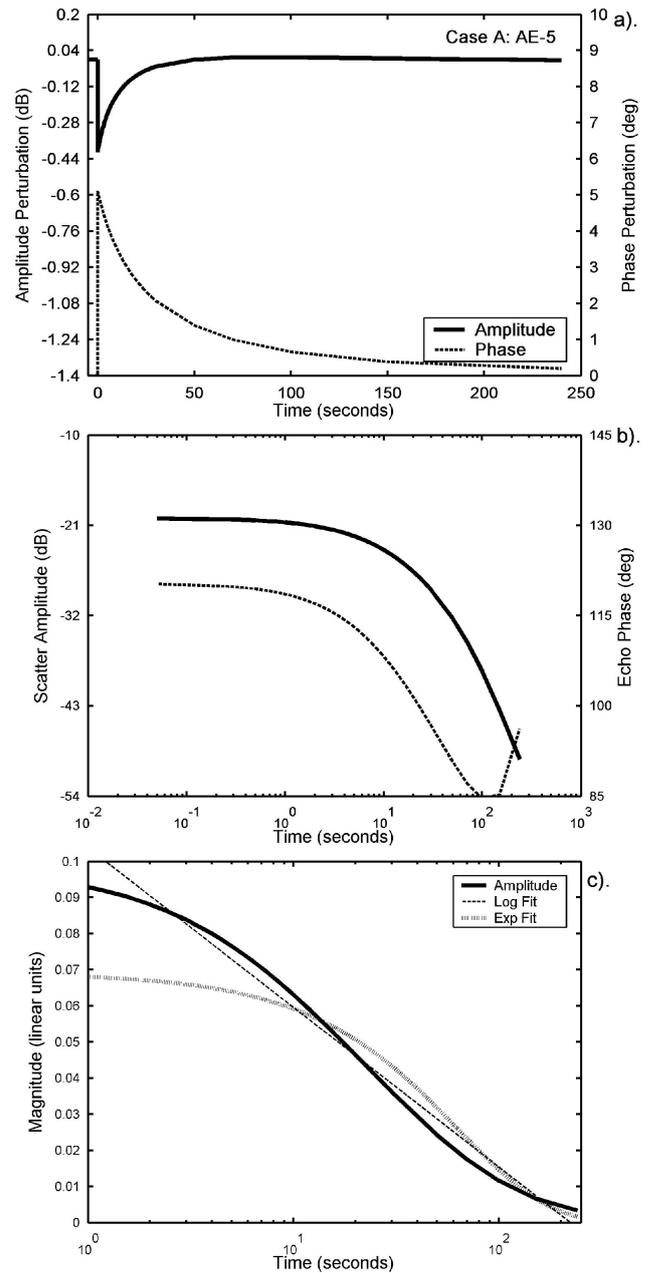


Fig. 4. VLF perturbation (“Trimpri”) calculated for NPM transmissions received at Faraday due to an ionospheric perturbation caused by a WEP burst with AE-5 energy spectra (case A) The top panel shows the perturbations in amplitude and phase, while the middle panel shows the calculated scatter magnitude and phase which produces this Trimpri. The lower panel shows the decaying scatter magnitude plotted on a logarithmic timescale after Dowden et al. (2001), with logarithmic (dashed) and exponential (dash-dot) fits.

of Fig. 4 the echo phase approaches $\sim 90^\circ$ as the perturbation decays, as noted above. The lower panel of Fig. 4 shows the decay of the scatter magnitude, M , on a logarithmic time scale. The scatter amplitude time decay features and their fit to the observations of Dowden et al. (2001) will be described in Sect. 4.1 below.

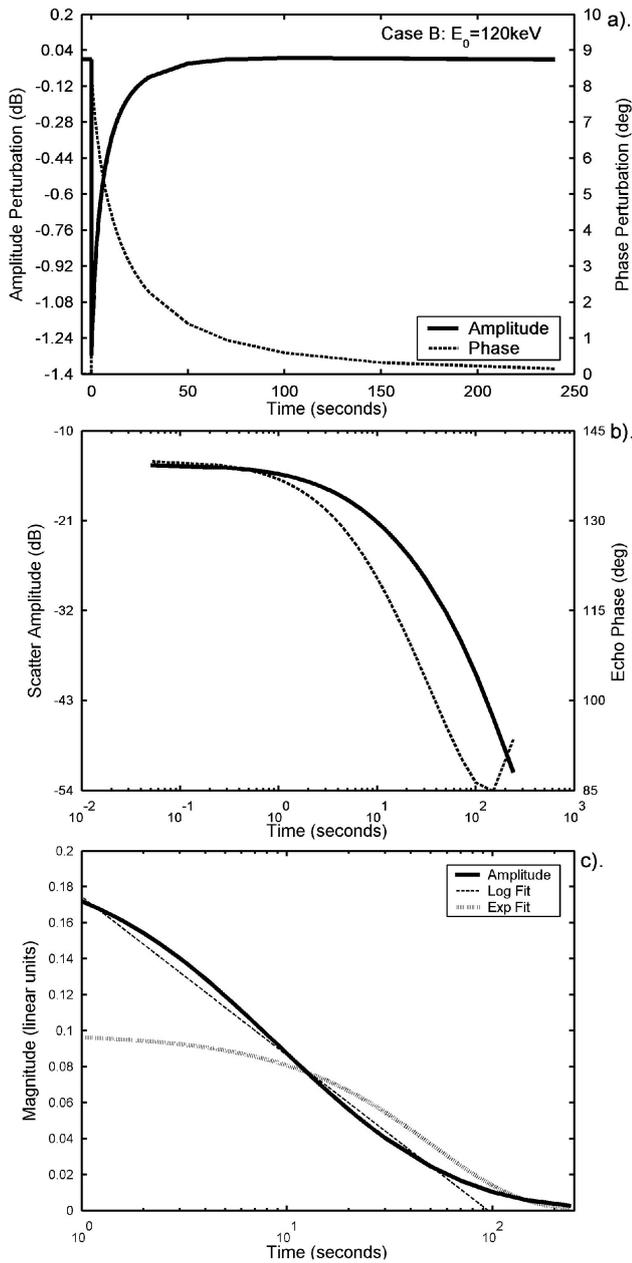


Fig. 5. VLF perturbation in the same form as that shown in Fig. 4, but in this case caused by a WEP burst with an e -folding energy spectra (case B).

Similarly, the calculated VLF perturbations created by the e -folding WEP burst (case B, Fig. 3b) are given in Fig. 5, in the same format as Fig. 4, for comparison with experimental observations.

3 Comparison with experimental observations

3.1 Temporal decay signature

The shapes of the decaying VLF perturbations shown in Figs. 4a and 5a resemble those observed in experimental data

(e.g. Dowden et al., 2001; Clilverd et al., 2002). For example, if one contrasts the WEP Trimpi presented in the upper panels of Figs. 10 and 11 of Dowden et al. (2001) with those in Figs. 4a and 5a one notes that the observed and calculated perturbations are positive (advancing) phase, negative (decreasing) amplitude, that the amplitude perturbation decays to its unperturbed state faster than the phase perturbation, and that the amplitude perturbation decays to the experimental noise level ($\sim 0.05 \text{ dB}$) in $\sim 25 \text{ s}$. Note that the Trimpi presented in the Dowden study were chosen on the grounds they are unusually large, and therefore well suited for examining the time decay profile of the perturbations. A comparison of the magnitudes with typical experimental data is discussed in the following section.

While the general shape of the calculated Trimpi are very similar to the experimental data, they are based on two cases of WEP fluxes with very different energy spectra. Do both cases produce the logarithmic decay profile observed experimentally? The decay signatures of the Trimpi amplitude and phase perturbations and the scatter magnitude are shown in Figs. 4 (case A) and 5 (case B). In the lower panels of these figures we undertake a comparison of our calculated decaying scatter magnitude with logarithmic (dashed) and exponential (dotted) fits. The experimentally observed decay of the scatter magnitude is best described through a logarithmic decay rather than an exponential (Dowden et al., 2001). This is also true for our calculated decaying scatter magnitude. Figure 4c shows that the logarithmic fit to the AE-5 (case A) WEP produced M over $t=1 \text{ s}$ to the end of the calculation period ($t=240 \text{ s}$), which is considerably better than the exponential fit. An indication of the “quality” of the fit can be found from the departure from unity of the square of the correlation coefficient, i.e. $1-r^2$. For the logarithmic fit $r^2=0.976$, but is only $r^2=0.895$ for the exponential case (i.e. ~ 4 times better fit); these values are similar to the WEP Trimpi event shown in Fig. 12 of the Dowden paper.

The situation is very similar when one considers the decaying scatter magnitude produced by the e -folding WEP burst (Fig. 5c). Again, the logarithmic fit is considerably better at describing the decay process ($r^2=0.972$) than the exponential ($r^2=0.83$). Clearly, there is no significant difference between the quality of the logarithmic fit between the WEP bursts using the AE-5 derived and e -folding energy spectra. Both lead to Trimpi which decay in a manner that is typical of experimentally observed events. As the decay process is primarily controlled by the atmospheric model used here we can take confidence in its confirmation of the logarithmic decay times. It has been argued that the WEP fluxes derived from the AE-5 Electron Model (case A) should be more typical than those described by an e -folding energy of $E_0=120 \text{ keV}$ (case B), as the latter is based on a single measurement (Rodger et al., 2003). We investigate this more closely in the section below.

3.2 Magnitude of perturbation

The AE-5 WEP burst (case A) shown in Fig. 4 produces a maximum Trimpi perturbation of -0.4 dB in amplitude and $+5^\circ$ in phase (-20 dB in scatter amplitude and 120° echo phase). In contrast, the higher WEP fluxes at larger energies in the e -folding case (case B; Fig. 5) result in a maximum Trimpi perturbation of -1.3 dB and 8.4° (-14 dB in scatter amplitude and 140° echo phase). Both are realistic in terms of the logarithmic decay times that they produce in the atmospheric model, as seen in the previous section. However, can we say anything about the likelihood of observing such Trimpi at any given time? Trimpi have been analyzed on the same path over an extended observation period (Clilverd et al., 1999). Trimpi with amplitude changes of >1 dB occurred on only 0.2% of occasions, while amplitudes of >0.4 dB were nearly 50 times more common. This is also true of the relative occurrences of $>8^\circ$ and $>5^\circ$ phase Trimpi (0.2% of occasions and 16 times more common).

The most commonly observed Trimpi magnitude is one that is just above the noise level. This is probably because of precipitation patch location, distribution of lightning magnitude, and changeable wave coupling into the plasmaspheric propagation path, thus altering the location and nature of the LIE. In the Clilverd et al. (1999) study there is no description of the location of the causative lightning discharge that lead to the Trimpi. Thus, our comparison with the Clilverd study above is likely to be influenced by enhanced occurrences of smaller Trimpi caused by positional variability at the lightning source (Clilverd et al., 2002). It is also possible that different lightning intensities will produce different WEP fluxes, as it is known that the WEP precipitation energy flux is proportional to the whistler wave equatorial magnetic field intensity, for most realistic wave fields (Inan et al., 1982). There are many more low current lightning events than large current ones (Uman, Fig. 7.8, 1987). This would fit the observations of many more small magnitude Trimpi perturbations than large ones. Evidence that confirms the association between lightning return current strength and Trimpi perturbation levels will be the subject of study in a future paper. As a result of these deliberations we can say that the Trimpi magnitudes given in this study suggest that the case A WEP electron spectra produce more typical Trimpi signatures than case B, given typical Trimpi observations and the typical precipitation flux employed.

3.3 Dependence upon LIE patch size

Experimental observations of Trimpi events observed on the signals from multiple Northern Hemisphere VLF transmitters received at multiple sites around the Antarctic Peninsula have shown that LIE patches are at least 1500 km along the major axis and perhaps 600 km along the minor axis (Clilverd et al., 2002). Observed Trimpi show high variability in amplitude and phase from event to event. Here we test to see if varying the size of the LIE patch can significantly influence the range of Trimpi perturbations seen or if the

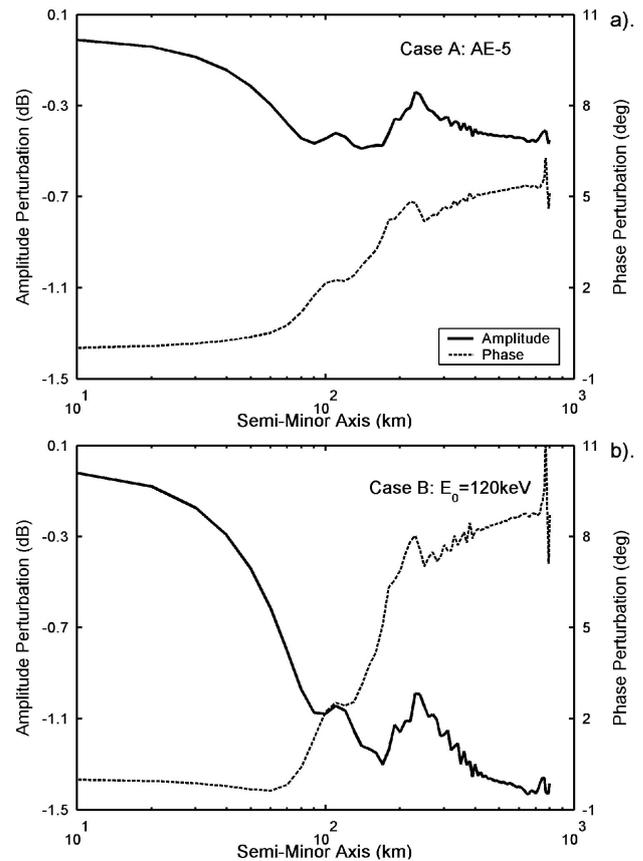


Fig. 6. Variation in the amplitude and phase of a Trimpi perturbation caused by LIEs with varying semi-minor axis sizes (a) LIE caused by a WEP burst with AE-5 energy spectra. (case A) (b) LIE caused by an e -folding energy spectra (case B).

variability is due to other factors such as precipitation flux levels. Figure 6 shows the calculated Trimpi magnitudes for varying LIE semi-minor axis sizes. In order to maintain the elliptical nature of the LIE-patch, the major axis was varied in proportion to the change in the minor-axis, maintaining the same ellipse ratio of 2.5, and the center point at the same location 800 km west of Faraday. The variation for very large LIEs is “computational noise”, due to multimodal effects, particularly due to those parts of the LIE patch very near the receiver point. This has been decreased by making calculations with a very fine grid, along with some averaging to decrease the significance of the noise remaining. As is clear, perturbation magnitudes increase steadily up to patch sizes of $\sim 200 \times 500$ km (semi-minor axis of ~ 100 km), after which there is little dependence upon the size of the patch. LIE patches produced by quasi-ducted or oblique whistlers propagating in ducts would be considerably larger than 200×500 km. Combining these results with the calculations presented in Clilverd et al. (Fig. 5, 2002) it appears that variations in the LIE patch size will lead to relatively small changes in Trimpi, and hence that the large variation in observed Trimpi is primarily due to changing WEP spectra and precipitation flux. The choice of LIE patch size will

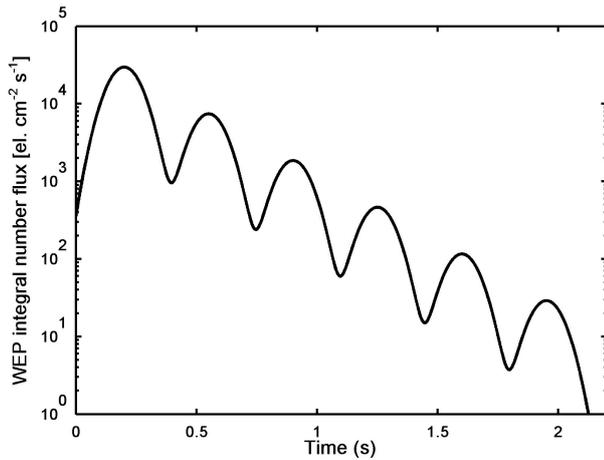


Fig. 7. Time dependence of the WEP electron flux used to simulate the decreasing train of precipitation pulses caused by the backscattering of precipitating electrons.

not strongly influence the findings of our study, or alter the calculations undertaken in the Clilverd paper.

3.4 Effect of pulsed WEP on Trimpi signature

In the earlier modeling we assumed that the WEP arrived during a single 0.2-s burst. However, in-situ satellite observations of WEP indicate that a portion of the precipitated electrons backscatter from the atmosphere and bounce repeatedly between the Northern and Southern Hemispheres, leading to a series of decreasing WEP pulses into both hemispheres over a period of ~ 2 –3 s (Voss et al., 1998). In order to investigate this effect, we have constructed a WEP burst made up of a train of pulses, roughly modeled on the time signature of the S81-1 satellite observations (Voss et al., Fig. 15, 1998). We estimate that each flux pulse was roughly half the magnitude of that occurring in the conjugate hemisphere, and therefore find the magnitude of the WEP pulse train (Fig. 7) by assuming that the ionospheric modification produced by the first pulse in the train will be half of that determined for the 0.2 s AE-5 WEP pulse (case A) in Sect. 2.2.

The time varying electron density modification has been calculated, as discussed in Sect. 2.3, using the time varying WEP pulse train. In this case, however, we allow for ionospheric relaxation to take place during the arrival of the ~ 2 -s pulse train, and determine $dN_e(z,t)$ with 0.05 s time resolution. The time-varying VLF perturbation calculated due to scattering from this ionospheric modification is shown in Fig. 8. Only the first 3 s of the perturbation are presented, as the pulsed nature of the WEP affects only the onset of the perturbation. Note that while additional WEP pulses arrive over ~ 2 s (Fig. 7), the VLF perturbation stops responding to production of extra ionisation after ~ 1.2 s into the onset phase. The ionospheric relaxation taking place between the bursts leads to a Trimpi perturbation that is slightly weaker in both phase and amplitude, such that the case A WEP spectra are even more likely to be typical.

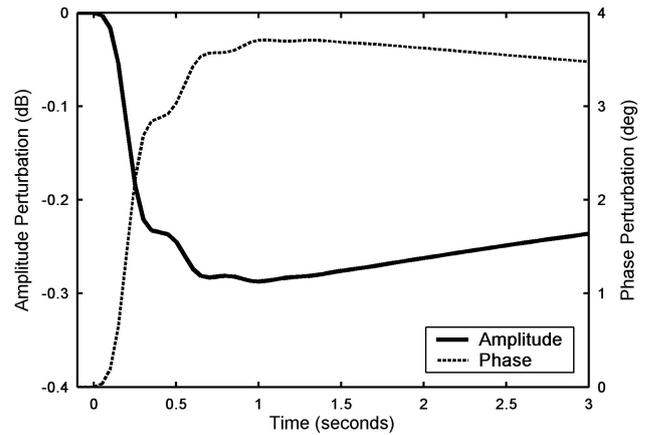


Fig. 8. Calculated onset of VLF perturbation due to the WEP pulse train shown in Fig. 7.

Experimental evidence for multi-burst WEP events has not been reported to date, although this may be because Trimpi observations with high enough time resolution measurements are rarely made. Nonetheless, the backscattering and bouncing of precipitating electrons is probably happening in the majority of WEP events, leaving an ionospheric signature which might be detectable in Trimpi perturbations. As there are a number of ionospheric processes which can lead to VLF perturbations (e.g. Rodger, 2003) similar to these “classic” Trimpi, this bursty onset might be a useful signature for discriminating WEP Trimpi from non-WEP Trimpi. The latter perturbations appear to be caused by much faster processes occurring inside and above thunderstorms, e.g. red sprites, elves, and sprite haloes (Rodger, 1999; Moore et al., 2003).

It should be noted that while each WEP event is made up of bursts, multi-hop whistlers are expected to precipitate significantly more electrons from the radiation belts than single hop whistlers.

3.5 Simulation of double Trimpi events

In order to test our modeling process we simulate the double Trimpi event shown in the top panel of Fig. 9. This was observed on the NPM signal at Faraday on 23 April 1994, shortly after 06:00 UT, and was selected as it is clearly defined and shows two events well above the noise. As seen in the figure, the initial NPM Trimpi occurs at 06:00:11 UT as a perturbation of -1.45 dB/ $+9.75^\circ$, followed at 06:00:17 UT by a Trimpi of -1.2 dB/ $+11.5^\circ$, i.e. the peak of the second amplitude perturbation is slightly less than the first Trimpi, while the peak in the phase perturbation is slightly larger than the first. The first of these Trimpi is probably associated with a cloud to ground lightning discharge detected by the National Lightning Detection Network (Cummins et al., 1998), occurring at about 06:00:10 UT (33.932° N, -74.986° E, peak return stroke current of 136.0 kA), and located near the region known to be associated with Trimpi observed in

the Antarctic Peninsula (Clilverd et al., 2002). There are three candidate NLDN strikes at about 06:00:16 UT which could have produced the second Trimpi. All are nearby to one another and have fairly large peak return stroke currents (31.155° N, -77.096° E, 44.5 kA), (31.177° N, -77.537° E, 88.7 kA) and (31.483° N, -77.504° E, 61.3 kA).

We simulate this situation by using WEP spectra to create a modified ionosphere as described above, centered 800 km west of Faraday station. The ionosphere is allowed to relax for 5 s as before, after which another WEP event causes additional ionisation to be added to that which existed from the first WEP. On the basis of the modified ionosphere the time varying Trimpi has been calculated and is shown in Fig. 9 for contrast with the experimental data.

As the first of these Trimpi are comparatively large when contrasted with typical events (Sect. 4.2), we employ the case B WEP spectra to model the first burst, using a mean precipitation energy flux of 2×10^{-3} ergs $\text{cm}^{-2} \text{s}^{-1}$, as before. The calculated VLF perturbation for this initial burst is shown in Fig. 9b. As the second Trimpi has a smaller amplitude perturbation than the first, one possibility is to use a smaller mean precipitation energy flux. We find that a WEP burst using the case B WEP spectra but a flux 65% lower than the first leads to a reasonable second Trimpi, as shown in Fig. 9b. The lower mean precipitation energy flux needed to simulate the second WEP burst is probably due to combination of the lower peak return stroke currents for the lightning which may be associated with this Trimpi (~ 30 – 65% lower), but also to less efficient coupling of the lightning energy through the ionosphere to produce the WEP itself. For example, it is likely that the second set of lightning is displaced further from the whistler duct entry point than the first stroke, and thus might be expected to have a proportionally smaller effect. The main features of the Trimpi decay are well reproduced in the calculated events, and the perturbations sizes are fairly similar. The primary difference is the decay time, where both the amplitude and phase perturbations decay away slightly faster in the experimental data than in the simulation. This may be an indication of the variability in ionospheric chemistry (Pasko and Inan, 1994), with the effective recombination coefficient at ionospheric heights showing high scatter (Gledhill, 1986). As the agreement between the form and decay of the double Trimpi event simulation and observed activity is high, we conclude that our process is sound.

As a contrast, we have also determined the Trimpi perturbation expected from a case B WEP burst followed by a case A WEP burst, using a mean precipitation energy flux of 2×10^{-3} ergs $\text{cm}^{-2} \text{s}^{-1}$ for both, and timed as above. This is shown in Fig. 9c. Again, the main features of the Trimpi decay are well reproduced, although the phase perturbation appears to be much longer lived. One could, however, argue again that the weaker lightning currents associated with the second Trimpi might have caused a weaker WEP energy flux. This shows that there is still significant parameter flexibility, such that one could choose an appropriate combination of spectra and energy flux and reproduce the double Trimpi

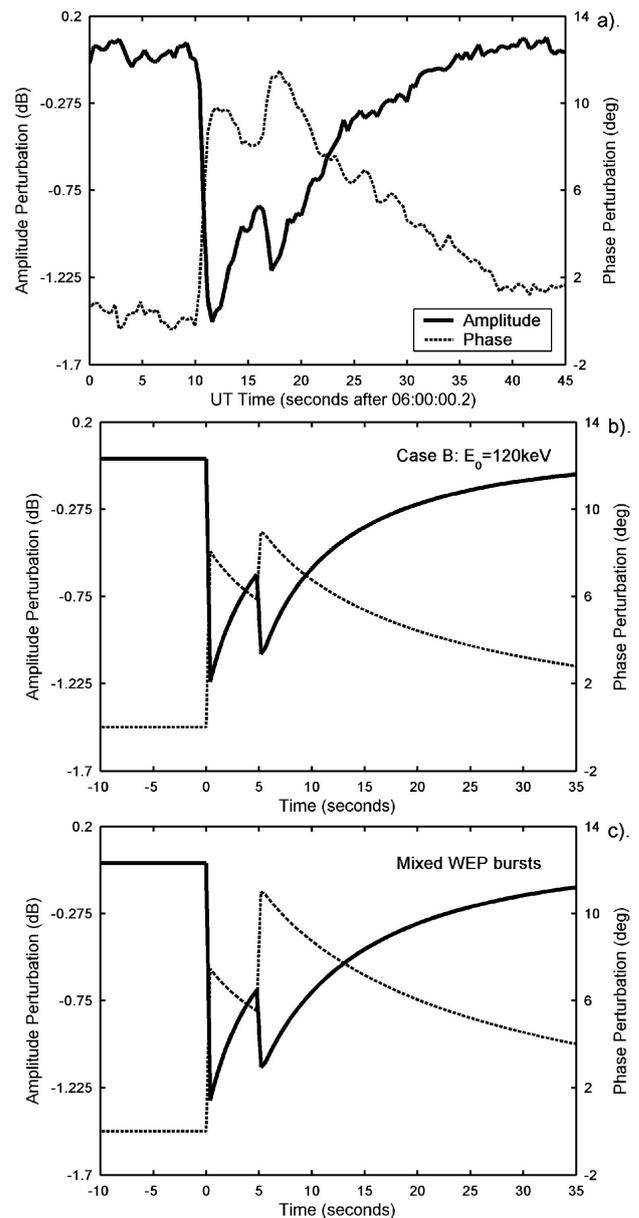


Fig. 9. Examination of the decay of two closely spaced WEP pulses producing a double Trimpi event. (a) Experimentally observed double Trimpi event, detected on NPM received at Faraday. (b) Calculated Trimpi perturbation caused by a case B WEP burst followed by another, smaller case B WEP burst. (c) As panel (b) but for a case B WEP burst followed by case A burst.

observed. An alternate approach for simulating this Trimpi event would be to use the case A WEP spectra to make the ionospheric modification that produces the first Trimpi, but with a larger precipitation energy flux to give larger perturbations than shown in Fig. 4. One can argue that the NLDN lightning discharge associated with this event is larger than typical lightning, and thus might be expected to produce a stronger WEP burst. However, at this stage it is not clear how lightning current should be related to Trimpi magnitude. In addition, while typical Trimpi are likely to be caused by

case A WEP bursts, our modeling shows that strong Trimperturbations may be due to WEP bursts with harder spectra or larger precipitation energy fluxes. Our double pulse modeling suggests that either effect is possible, i.e. that the variation in Trimperturbations may be due to changing WEP spectra or changing burst magnitude. However, given the short time separation between the double Trimperturbations shown in Fig. 9, it seems most likely that this is primarily due to changing precipitation energy flux, possibly driven by differing lightning currents. It is less likely that the WEP spectra will change from near-case A to near-case B in ~ 5 s, even though both spectra are valid in general. Thus we conclude that the double Trimperturbations event is strongly influenced by changing precipitation energy flux, despite nearly co-located lightning sources. The factor of three difference in flux levels observed in events separated by only 5–6 s requires further examination.

4 Summary and conclusions

In this paper we have undertaken calculations of the Trimperturbations and contrasted these with experimental observations so as to better understand the WEP fluxes producing the ionospheric modification, characterized through experimental WEP observations, and thus confirm the estimates of the likely impact of the WEP on the radiation belts.

It has been argued that the WEP fluxes derived from the AE-5 Electron Model (case A) should be more typical than those described by an e -folding energy of $E_0=120$ keV (case B) (Rodger et al., 2003), as the latter is based on a single measurement, albeit an in-situ observation which has been subject to significant analysis (Inan et al., 1989; Voss et al., 1998). The Rodger study noted that the selection of WEP spectra and the value of typical precipitation magnitude lead to a significant source of uncertainty in the calculated lifetime of radiation belt electrons. In this study we find that typical values employed by Rodger et al. (2003), based on the AE-5 electron model, appear reasonable, producing Trimperturbations of typical magnitudes. Given that Trimperturbations rates were used to determine the WEP rate in the Rodger study, the radiation belt lifetimes given in this study should be a reasonable upper bound for WEP driven lifetimes. As such, the conclusions of these authors as to the relative importance of WEP losses appear sound. It should be noted that this finding is limited by the limited range of WEP spectra which have been considered. While other WEP spectra might be as good, or better, both cases considered here are based on experimental observations of radiation belt particles and have been used in previous studies on WEP losses and ionospheric modifications.

In addition, it appears that calculated Trimperturbations are only weakly dependent upon the LIE patch size used to describe the WEP produced ionospheric modification. Variations in the patch size will lead to relatively small changes in Trimperturbations, and hence that the large variation in observed Trimperturbations is primarily due to either changing WEP spec-

tra or precipitation flux. However, when considering Trimperturbations that are closely spaced in time, it seems more likely that the variation in perturbation magnitude is driven by changes in the strength of the precipitation flux. Large Trimperturbations occurring repeatedly over a long time period are likely to be due to the radiation belt having a harder spectra than typical conditions. Variability due to lightning current and location will also play a major role.

Finally, by considering the onset of Trimperturbations it has been shown that high time resolution measurements may provide a useful signature for discriminating WEP Trimperturbations from non-WEP Trimperturbations, by exploiting the pulsed nature of the WEP arrival.

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